

***Did oceanic biogenic methane cycling
regulate the evolution of Early Earth
atmospheric chemistry?***



	Eonothem Eon	Erathem Era	System Period	Age Ma	GSSP GSSA
Precambrian	Proterozoic	Neo-proterozoic	Ediacaran	542	
			Cryogenian	~635	🔑
			Tonian	850	🔄
		Meso-proterozoic	Stenian	1000	🔄
			Ectasian	1200	🔄
			Calymmian	1400	🔄
		Paleo-proterozoic		1600	🔄
			Statherian	1800	🔄
			Orosirian	2050	🔄
			Rhyacian	2300	🔄
			Siderian	2500	🔄
	Archean	Neoarchean		2500	🔄
		Mesoarchean		2800	🔄
		Paleoarchean		3200	🔄
		Eoarchean		3600	🔄
	Hadean (informal)			4000	
				~4600	



- Ocean formation
- **Oxidation of the environment**
Ocean and Atmosphere
- Continental crustal growth
- **Development of life**

~ Half of the geological record

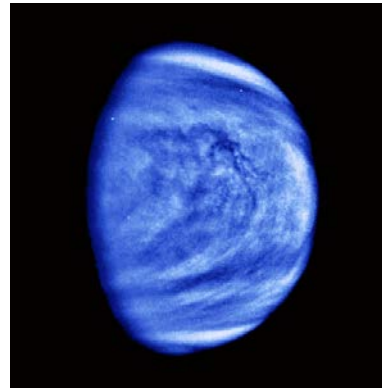
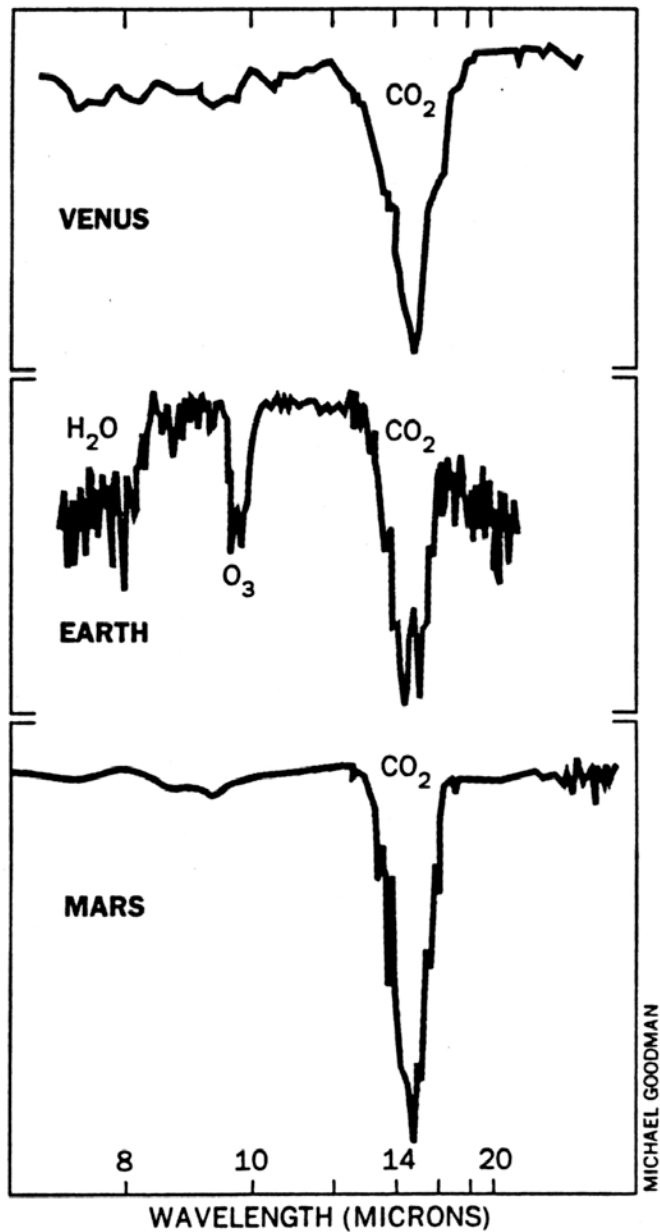
	Eonothem Eon	Erathem Era	System Period	Age Ma	GSSP GSSA
Precambrian	Proterozoic	Neo-proterozoic	Ediacaran	542	
			Cryogenian	~635	🔧
			Tonian	850	🔧
		Meso-proterozoic	Stenian	1000	🔧
			Ectasian	1200	🔧
			Calymmian	1400	🔧
		Paleo-proterozoic	Statherian	1600	🔧
			Orosirian	1800	🔧
			Rhyacian	2050	🔧
			Siderian	2300	🔧
	Archean	Neoarchean		2500	🔧
		Mesoarchean		2800	🔧
		Paleoarchean		3200	🔧
		Eoarchean		3600	🔧
	Hadean (informal)			4000	
				~4600	



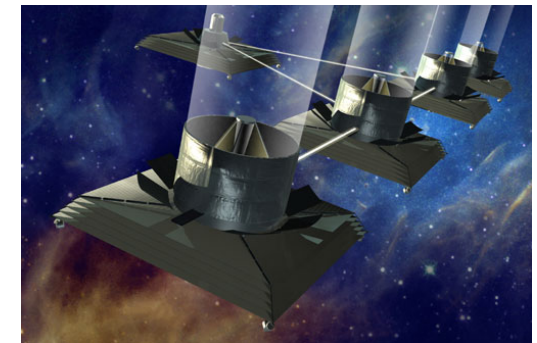
Poorly archived, few sediments, metamorphosed at $T > 200^{\circ}\text{C}$

Early life and its environment remains
poorly characterized

~ Half of the geological record



Thermal IR Spectra using TPF



R. Hanel, Goddard
Space Flight Center

A decade of multiple sulfur isotope studies bring new insights in the timing of the Earth system oxygenation



Atmospheric Influence of Earth's Earliest Sulfur Cycle

James Farquhar, *et al.*

Science **289**, 756 (2000);

DOI: [10.1126/science.289.5480.756](https://doi.org/10.1126/science.289.5480.756)

Theory of Mass dependent/Independent Fractionation of sulfur isotopes

4 isotopes :

^{32}S : 95,02%

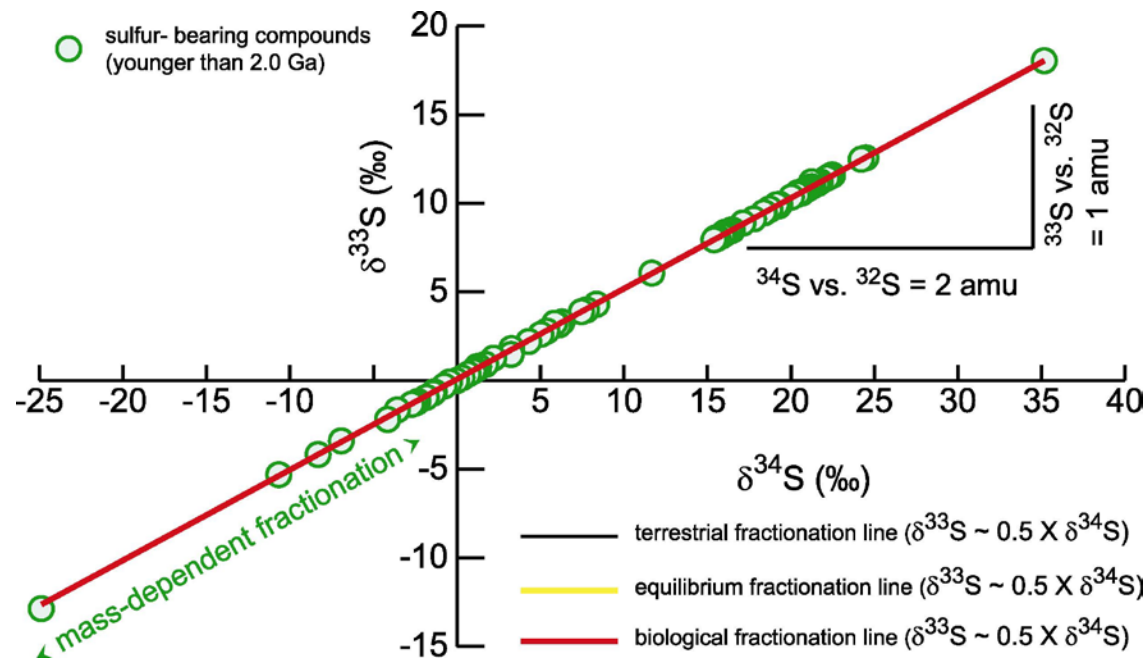
^{33}S : 0,75 %

^{34}S : 4,21%

^{36}S : 0,02%

$$\text{Mass } (^{34}\text{S}) - \text{Mass } (^{32}\text{S}) = 2 \times \text{Mass } (^{33}\text{S}) - \text{Mass } (^{32}\text{S})$$

← isotopic fractionation depends on mass ...



Prediction from theory:

$$\delta^{34}\text{S} = 2 \times \delta^{33}\text{S}$$

Theory of Mass dependent/Independent Fractionation of sulfur isotopes

4 isotopes :

^{32}S : 95,02%

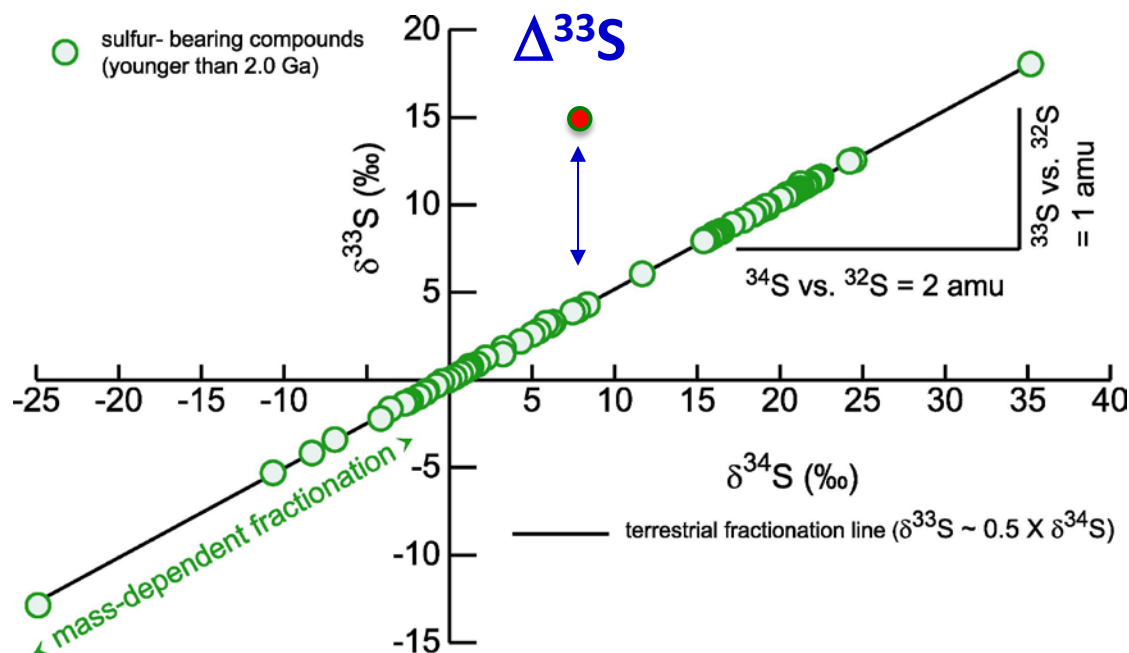
^{33}S : 0,75 %

^{34}S : 4,21%

^{36}S : 0,02%

$\text{Mass } (^{34}\text{S}) - \text{Mass } (^{32}\text{S}) \neq 2 \times \text{Mass } (^{33}\text{S}) - \text{Mass } (^{32}\text{S})$

← isotopic fractionation independent on mass ...



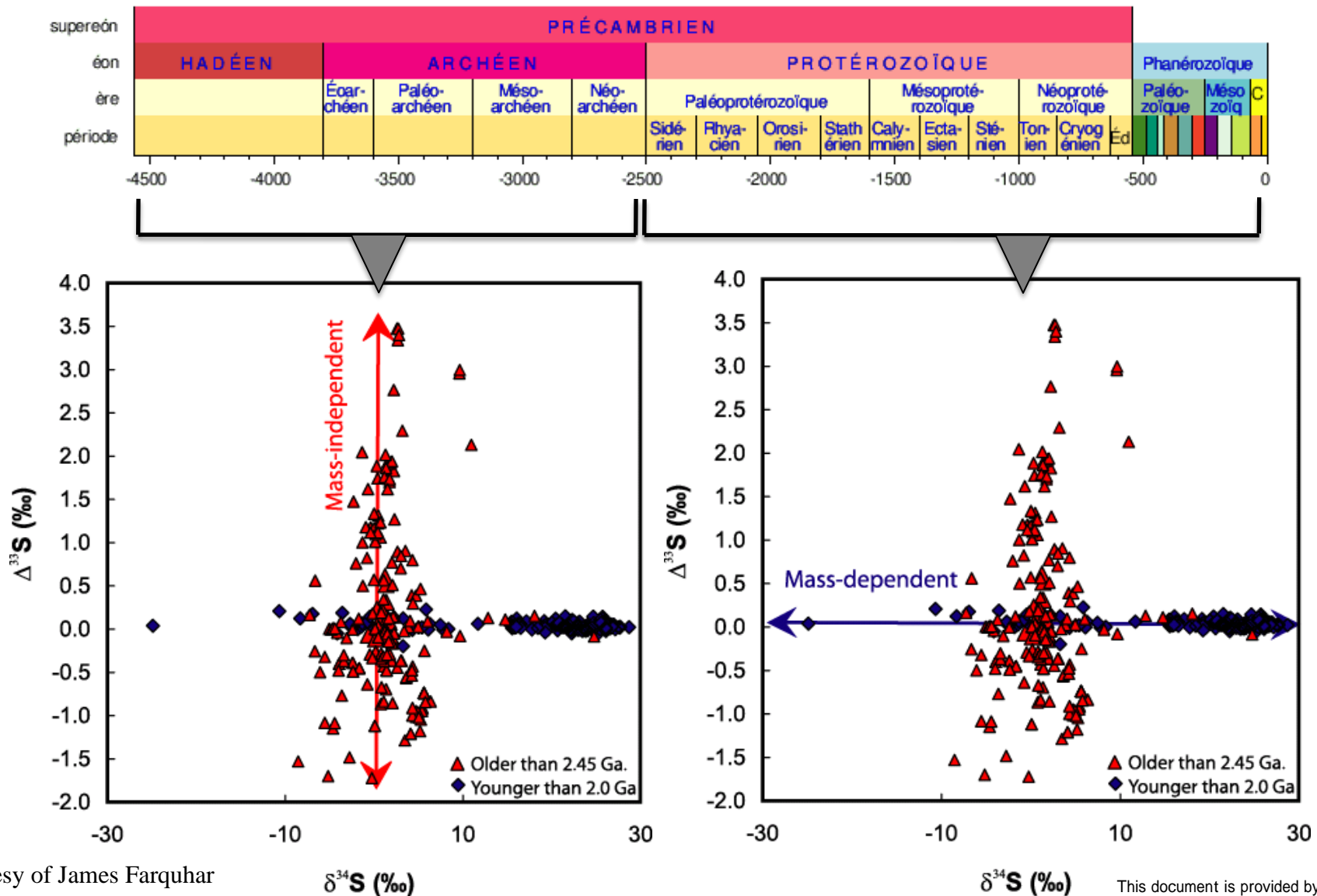
Prediction from theory:

$$\delta^{34}\text{S} \neq 2 \times \delta^{33}\text{S}$$

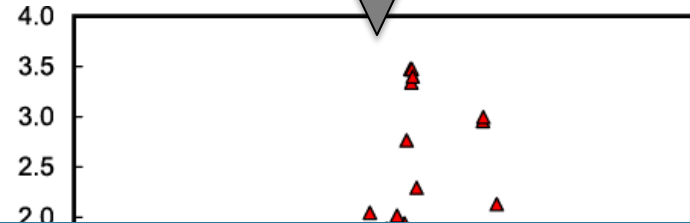
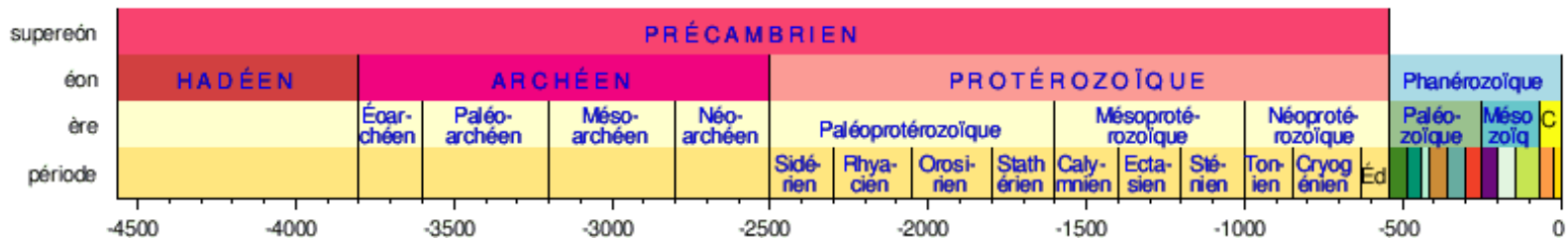
$$\leftarrow \Delta^{33}\text{S} \approx \delta^{33}\text{S} - 0.515 \delta^{34}\text{S}$$

deviation with regard to the
terrestrial fractionation line =
(MIFs)

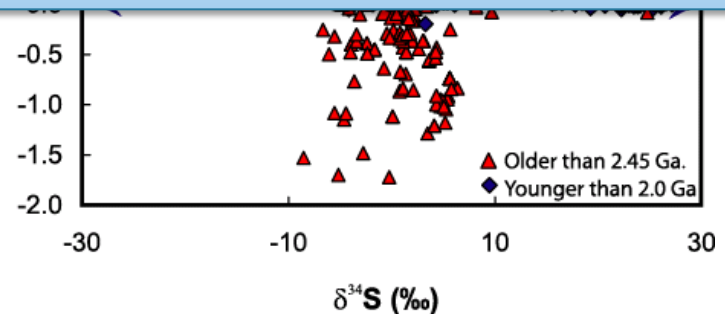
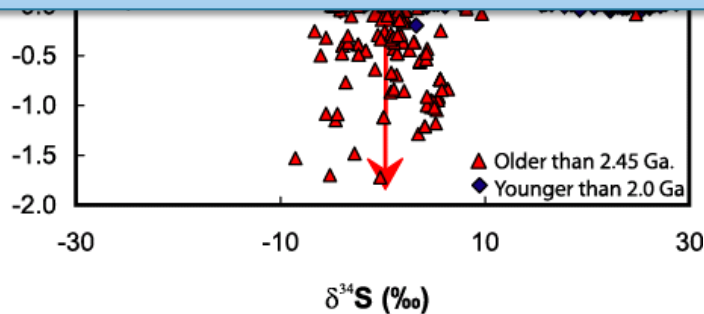
Geological record



Geological record

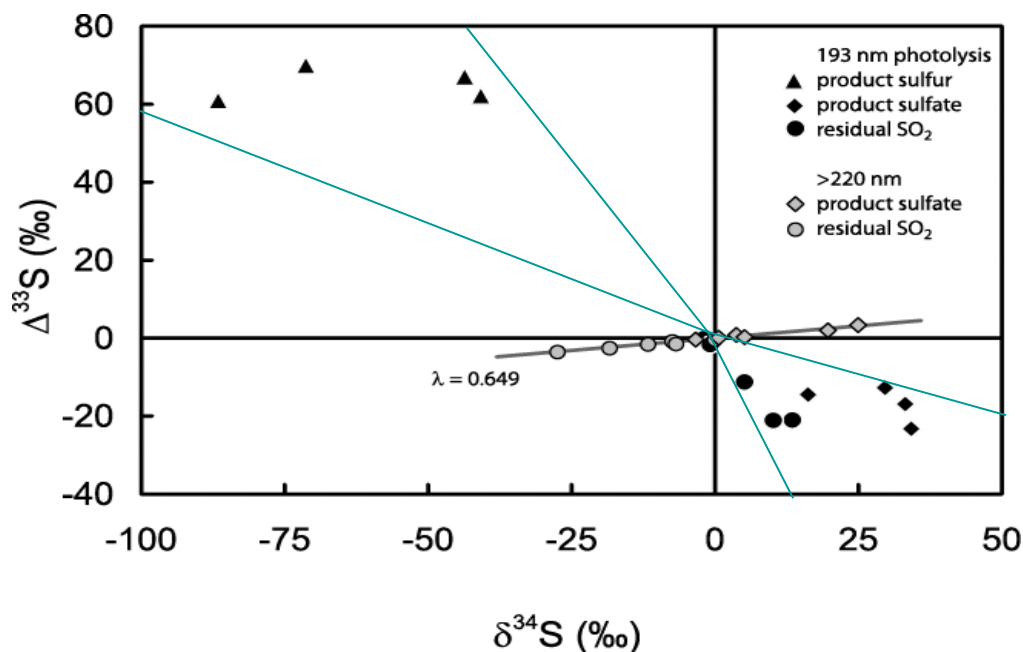


Mechanisms to produce and preserve MIF-S in the
> 2.45Ga sedimentary record ?



➤ MIF-S are produced in lab experiment during UV photolysis of SO_2

Farquhar et al, 2000



Photolytic reaction of S in anoxic atmosphere

Without O_2 (O_3)



Atmospheric Influence of Earth's Earliest Sulfur Cycle

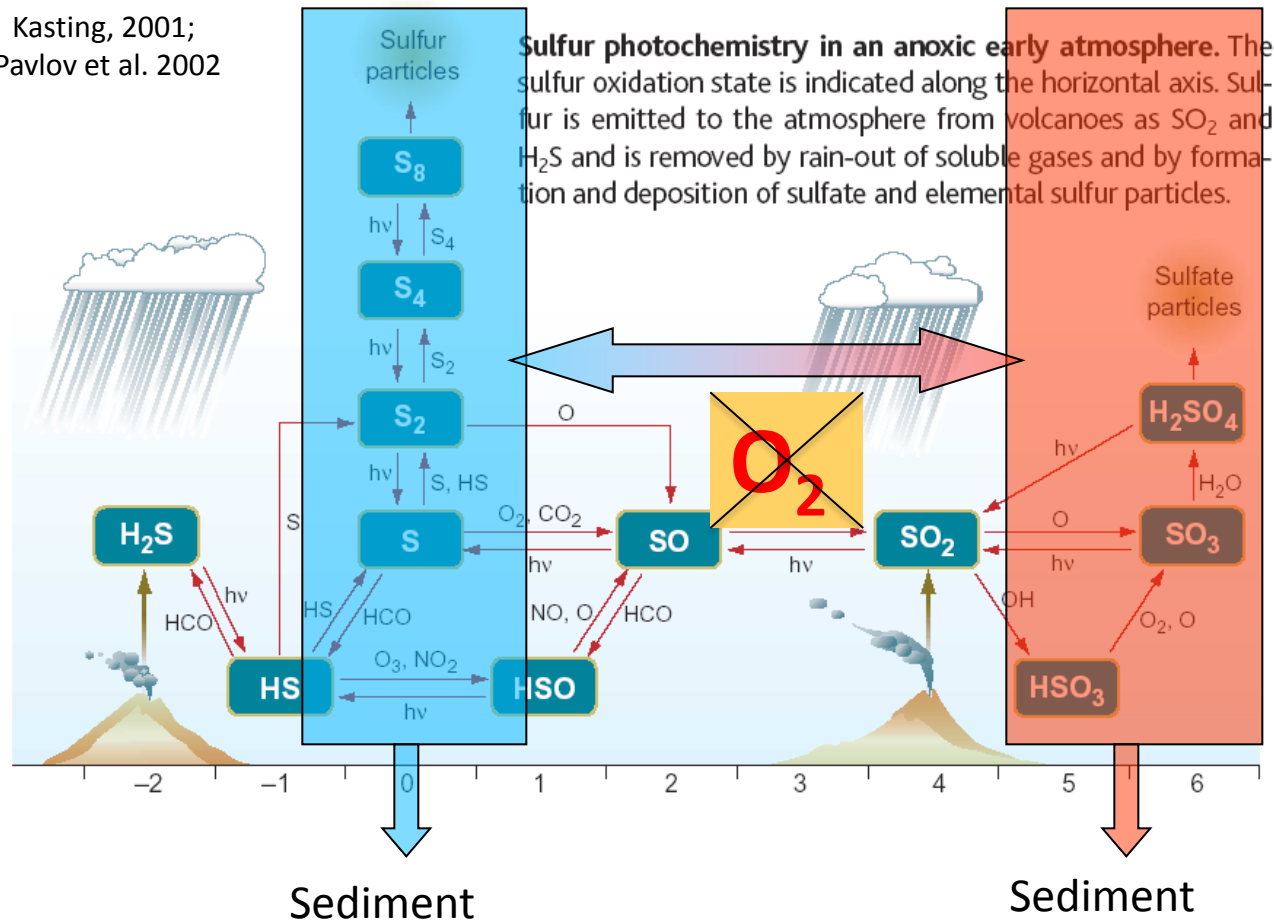
James Farquhar, *et al.*

Science **289**, 756 (2000);

DOI: 10.1126/science.289.5480.756

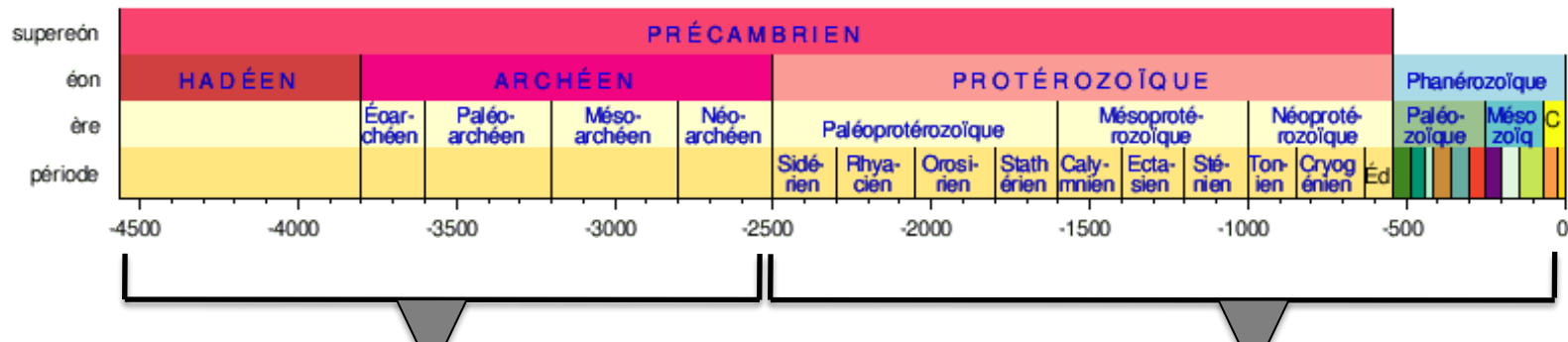
Preservation in the sedimentary record?

Kasting, 2001;
Pavlov et al. 2002

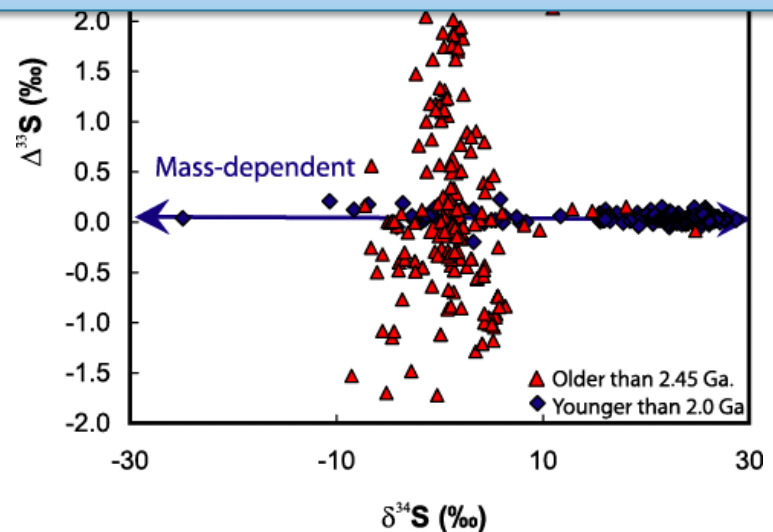
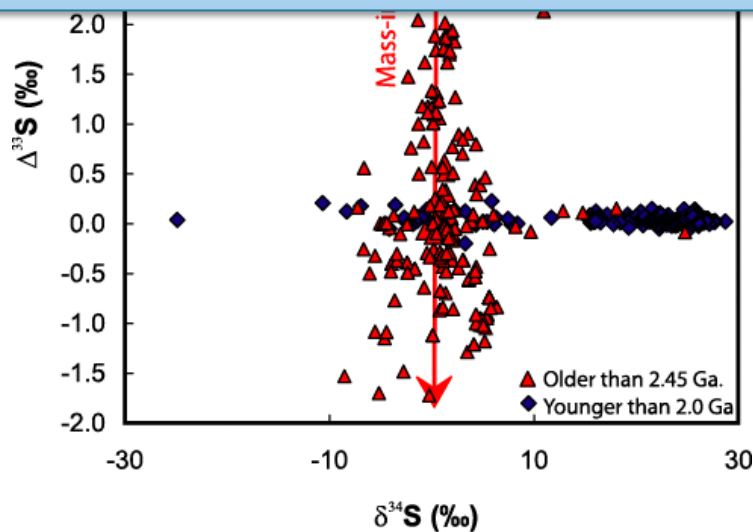


- Two exit channels of sulfur species is required and imply **low O_2 ($< 10^{-5}$ PAL)** in the Archean environment

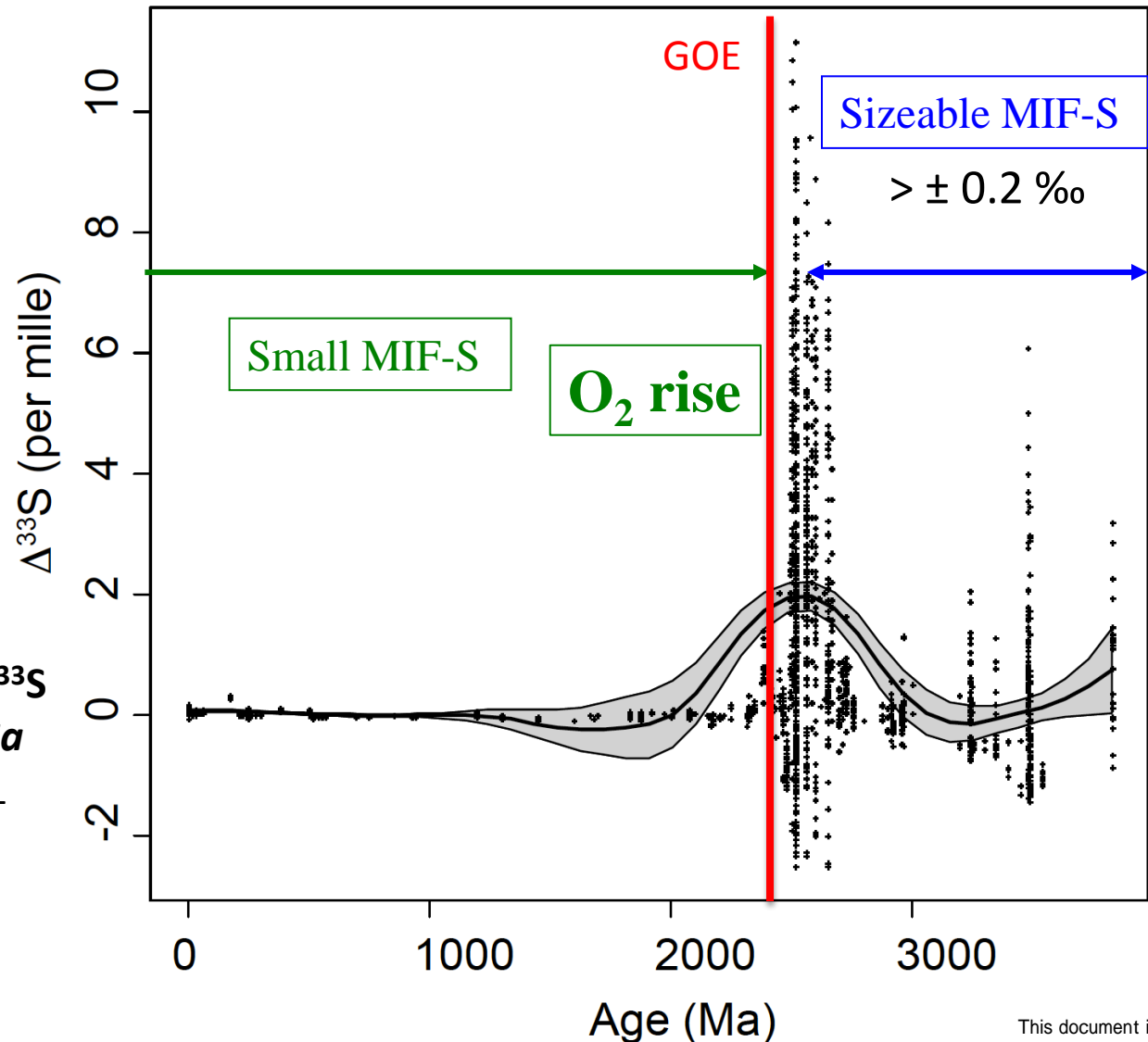
Geologic record



The transition at ~ 2.4 Ga are interpreted by most of the authors as the GOE (Great Oxidation Event)

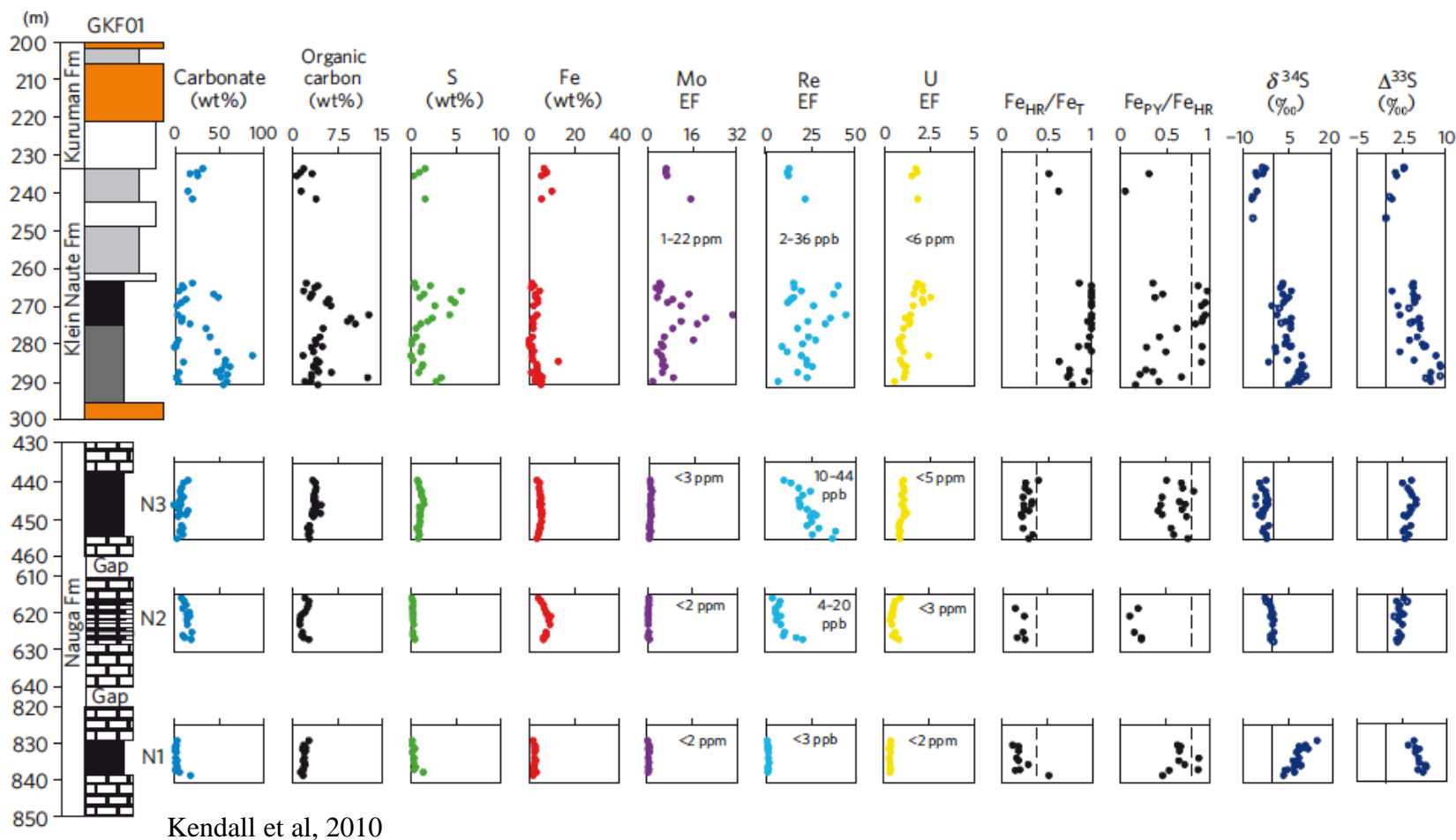


.... But in details there is some secular variations in the MIF-S record during the Precambrian ...



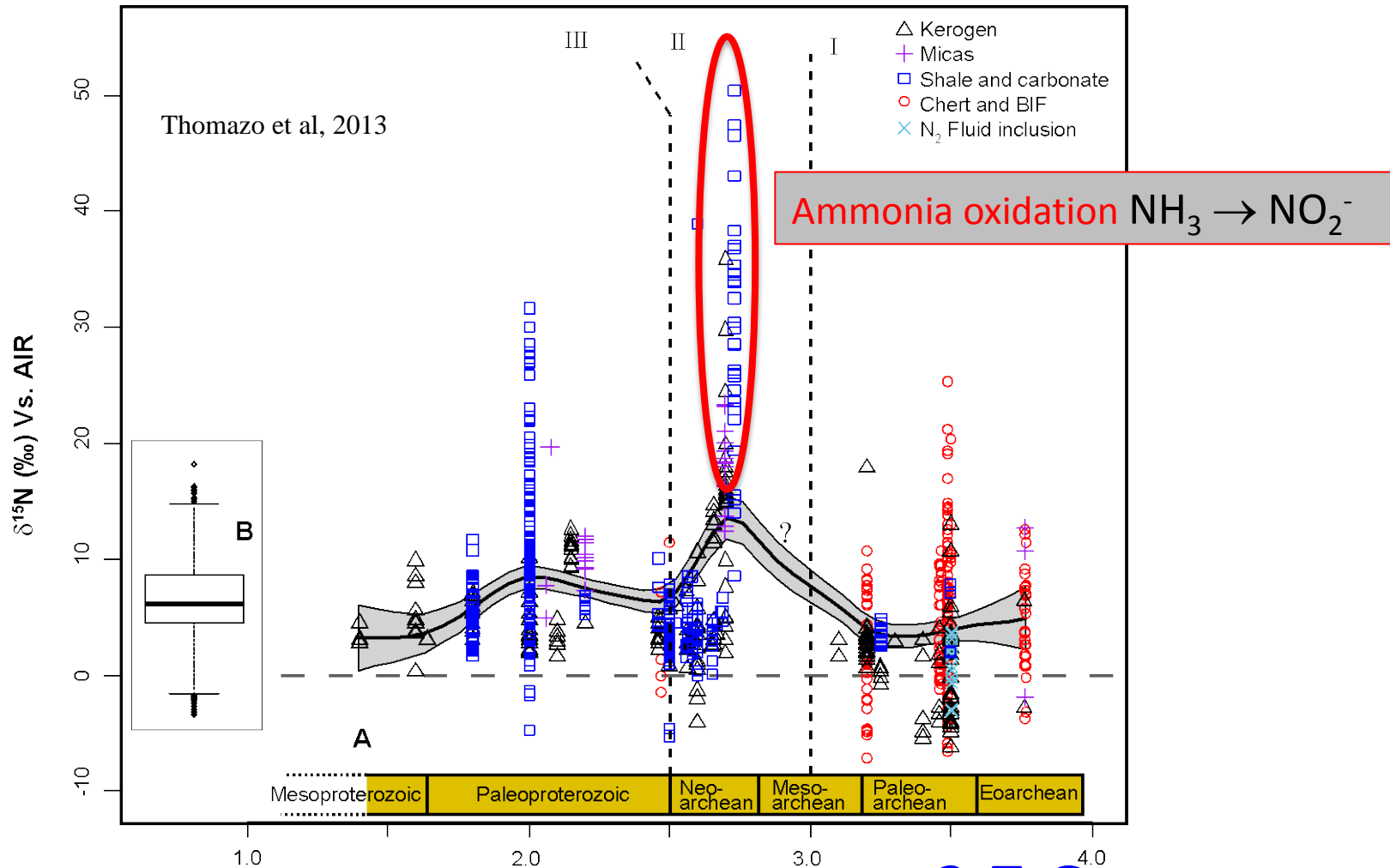
NB: dampen down $\Delta^{33}\text{S}$
between 3.1 to 2.7 Ga
(i.e. value between -1
to +1 ‰)

.... But in details there is some secular variations in the MIF-S record during the Precambrian ... AND some evidences for a « *whiff of oxygen* » before the GOE



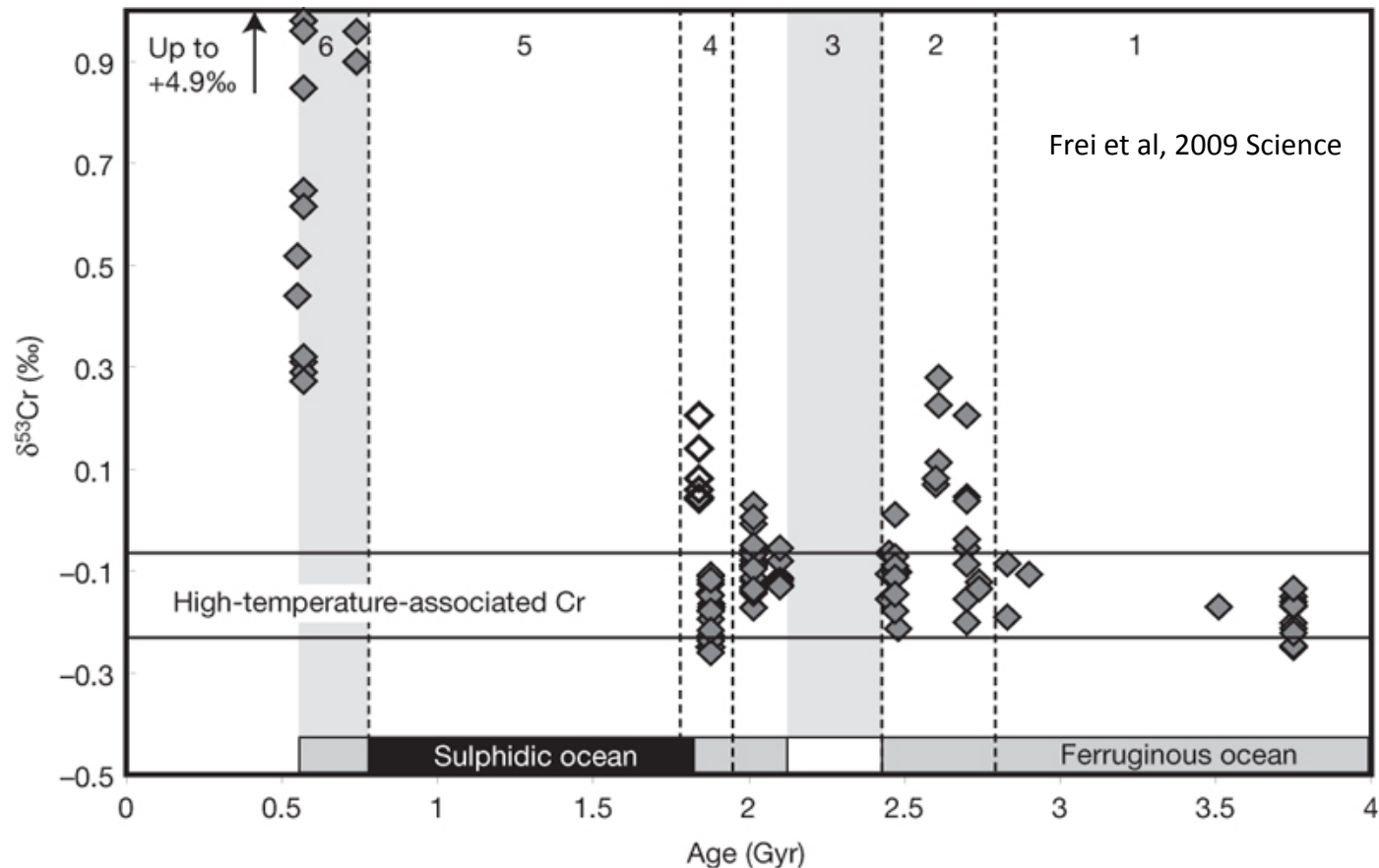
➤ Re enrichment and iron speciation point to oxidative weathering at around **2.6 Ga**

.... But in details there is some secular variations in the MIF-S record during the Precambrian ... AND some evidences for a « *whiff of oxygen* » before the GOE



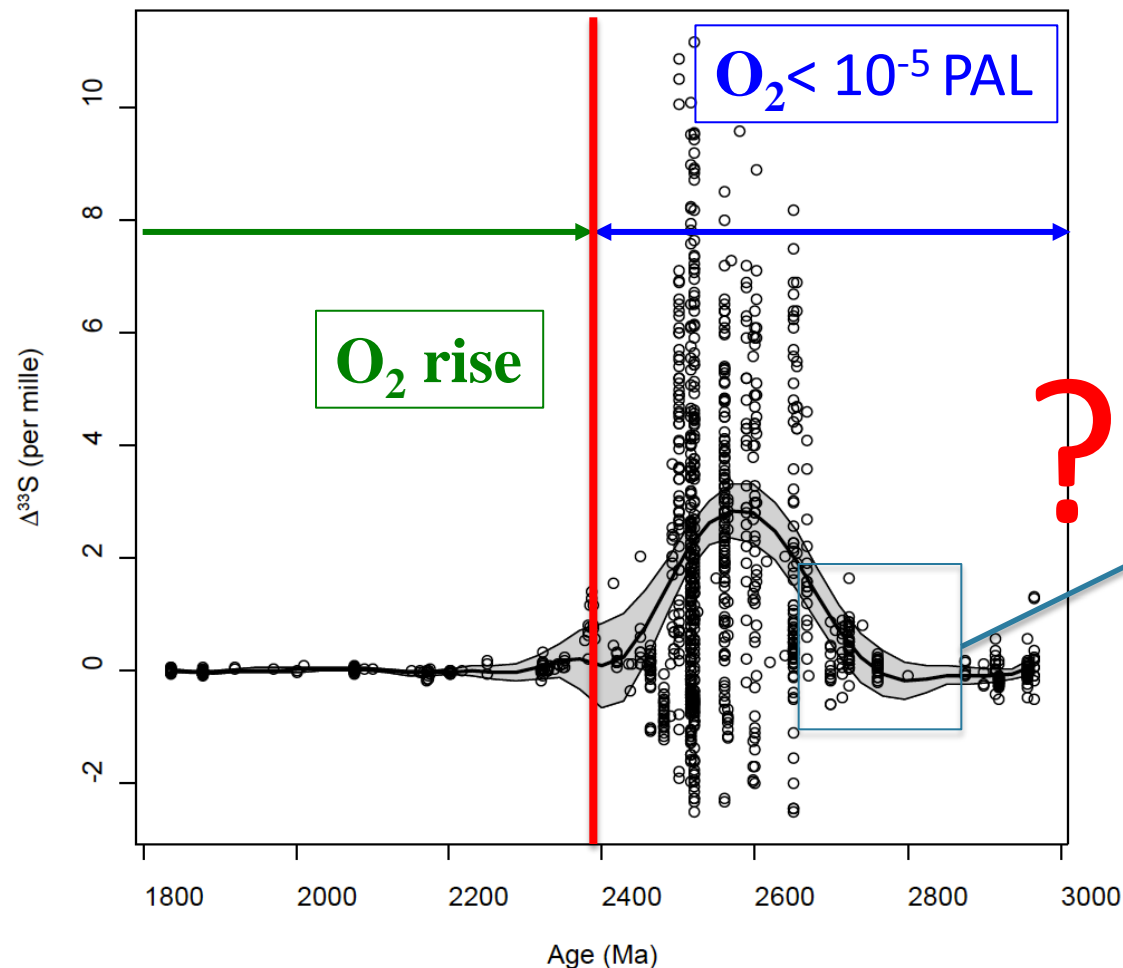
➤ The oxidative part of the nitrogen cycle is recorded SINCE 2.7 Ga

.... But in details there is some secular variations in the MIF-S record during the Precambrian ... AND some evidences for a « *whiff of oxygen* » before the GOE

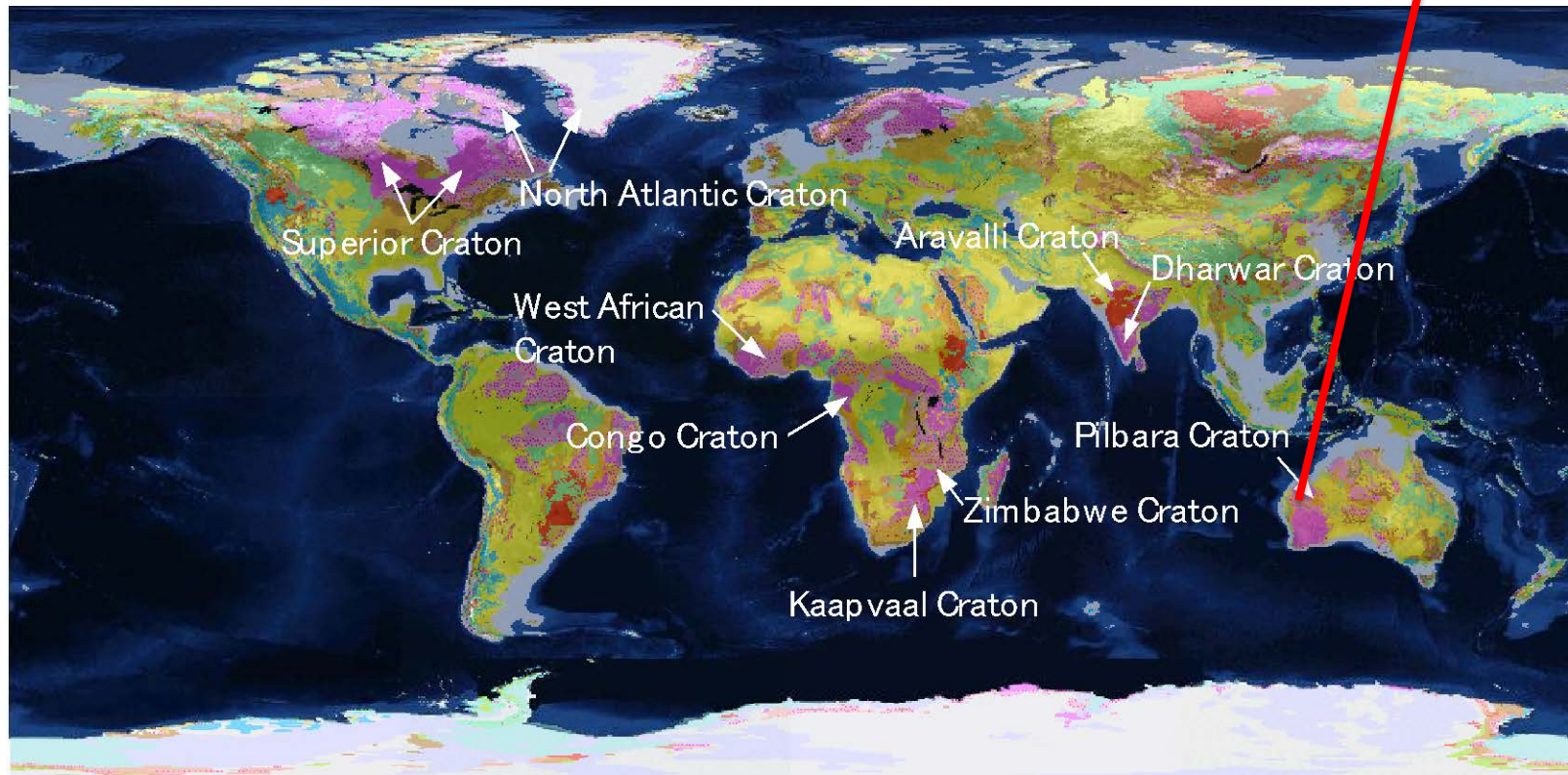


➤ accumulation of Cr(VI) in ocean surface waters **~2.8 to 2.6 Gyr ago**

How to council MIF-S secular variations (magnitude and sign) and a transient elevation in atmospheric and surface ocean oxygenation as early as 2.7 Ga ?



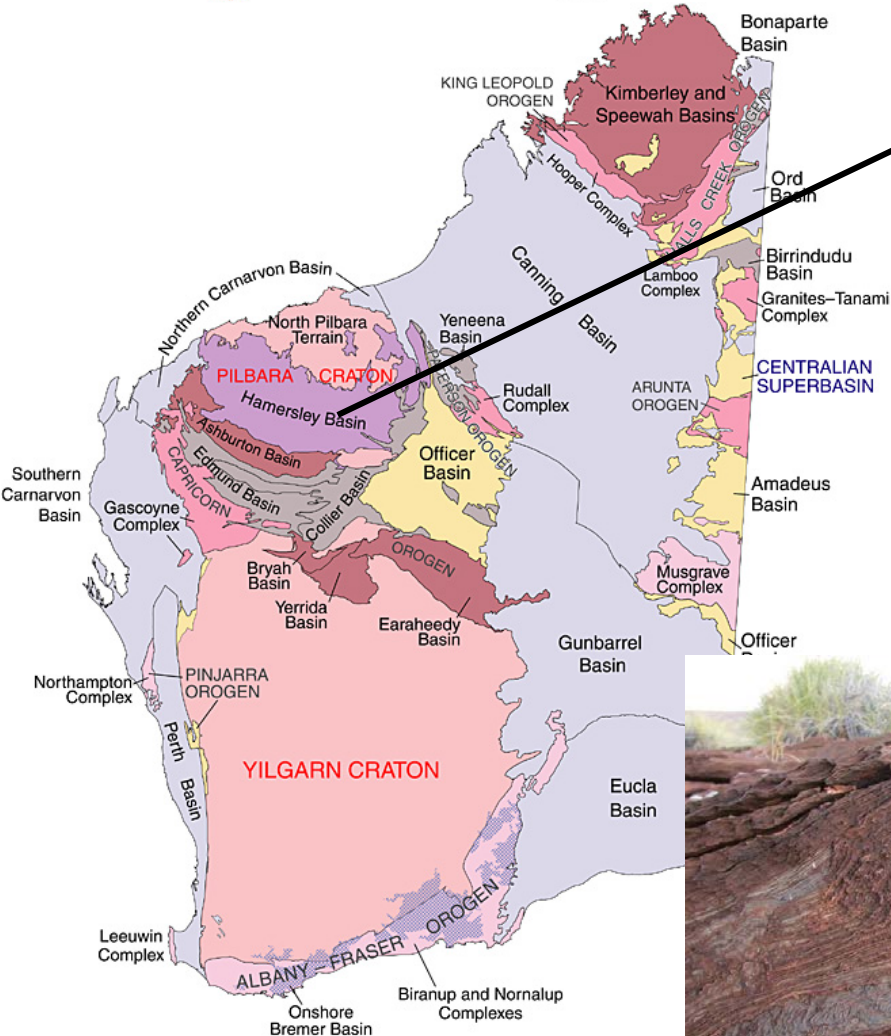
1/ Tumbiana Fm. (Pilbara Craton, Australia)



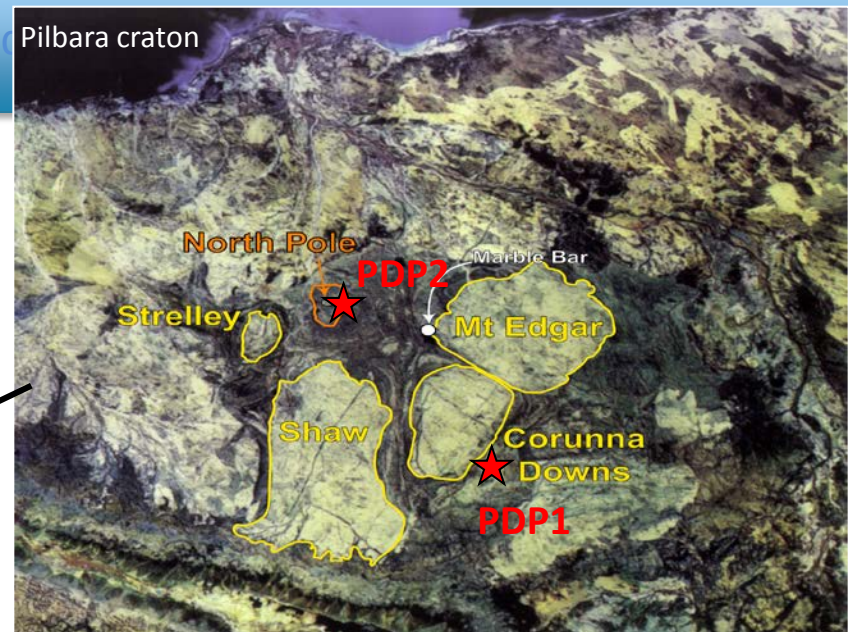
Onshore tectonic units of Western Australia



Department of
Mineral and Petroleum Resources



Pilbara craton



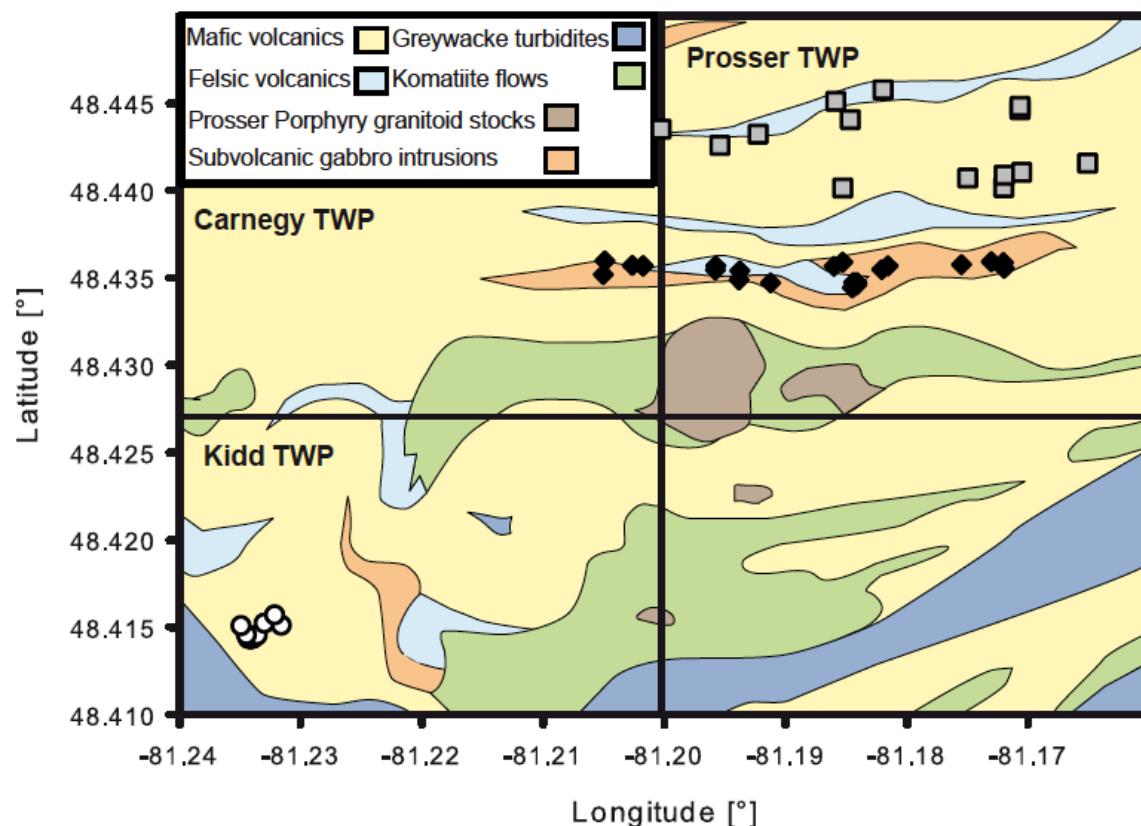
- PDP 1 Drill core (2004)
- 2.73 Ga Tumbiana Formation
- Shallow marine environment
- Subgreenschist facies



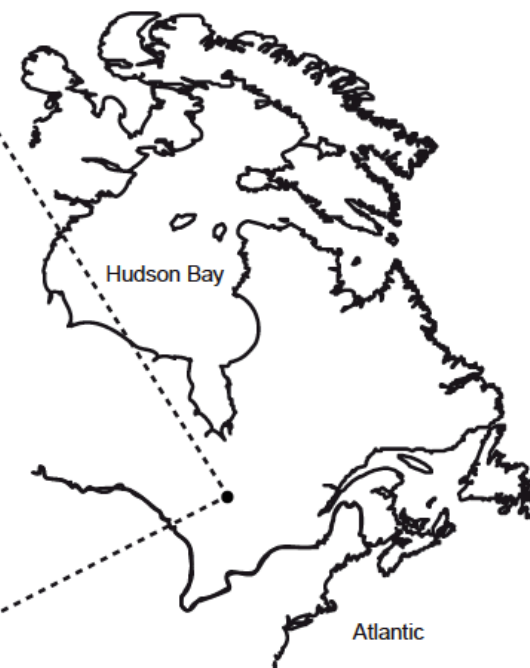
2/ Kidd Creek Fm. (Abitibi Subprovince, Canada)

1/ Tumbiana Fm. (Pilbara Craton, Australia)





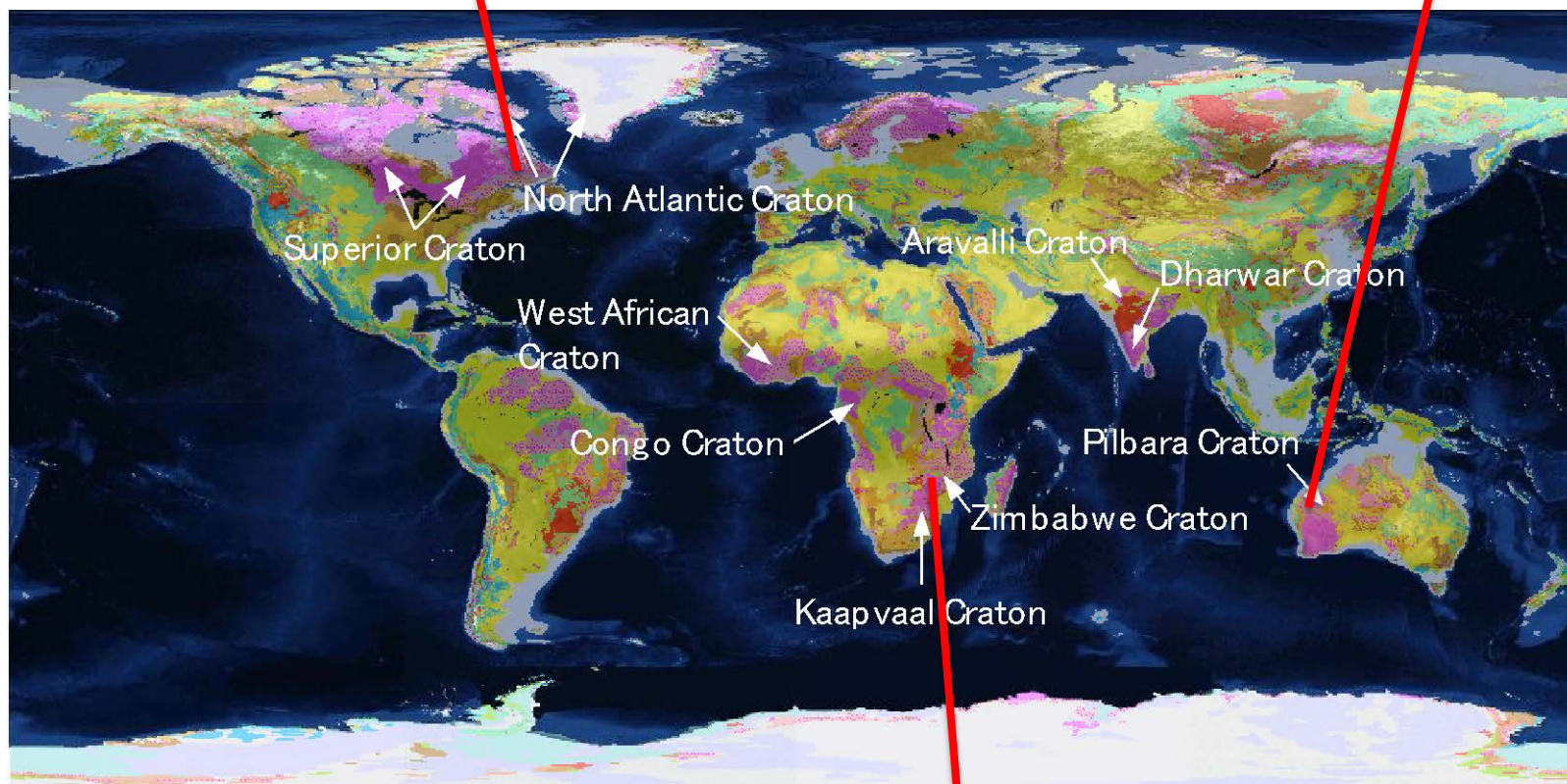
Kurzweill et al, 2013



- Outcrop samples (sedimentary argillite layers in between felsic volcanics)
- [2.71 Ga](#) for the sedimentary rocks of the Kidd Creek area
- Deposition in biologically active areas close to hydrothermal vents (Wellmer et al., 1999)
- Hydrothermal alteration below 250°C (Hannington et al., 1999b)
- The metamorphic grade of the Kidd Creek is sub-greenschist to lower greenschist

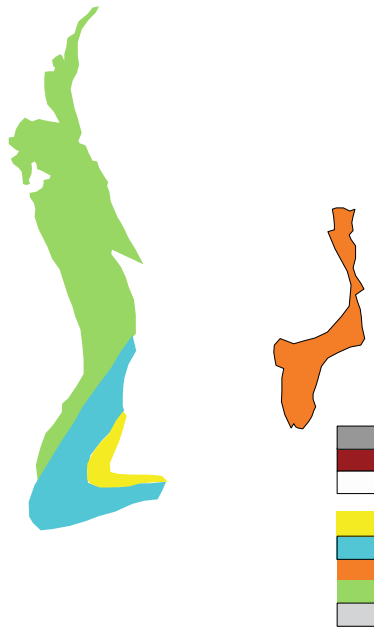
2/ Kidd Creek Fm. (Abitibi Subprovince, Canada)

1/ Tumbiana Fm. (Pilbara Craton, Australia)



3/ Cheshire Fm. (Ngezi Group, Zimbabwe)

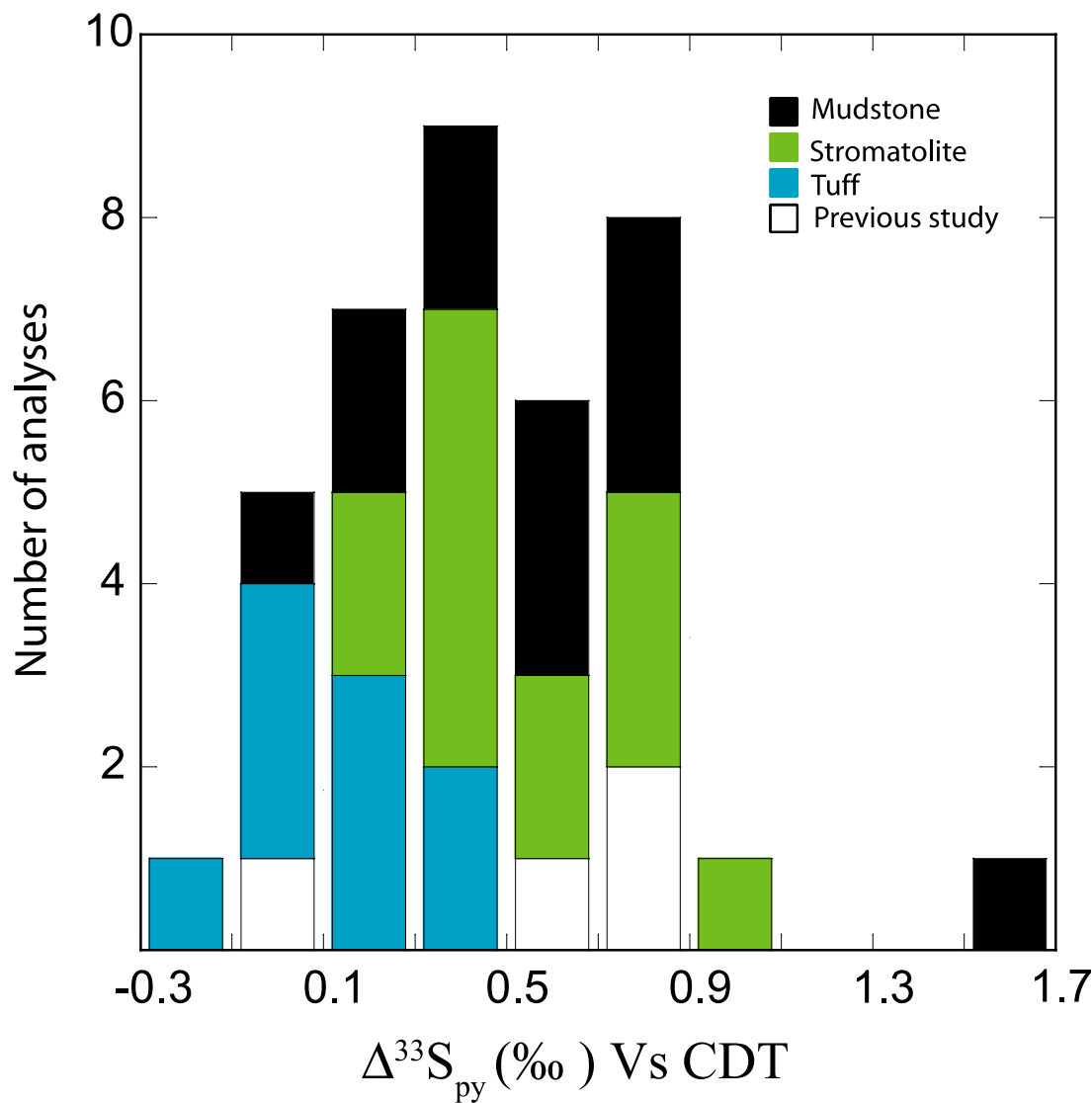
Thomazo et al, 2013



- Nercmar core and Saskmar holes (80')
- 2.65 Ga Cheshire Formation
- Shallow marine environment
- Samples are unstrained and in the subgreenschist facies

1/ 2.73 Ga Tumbiana Formation

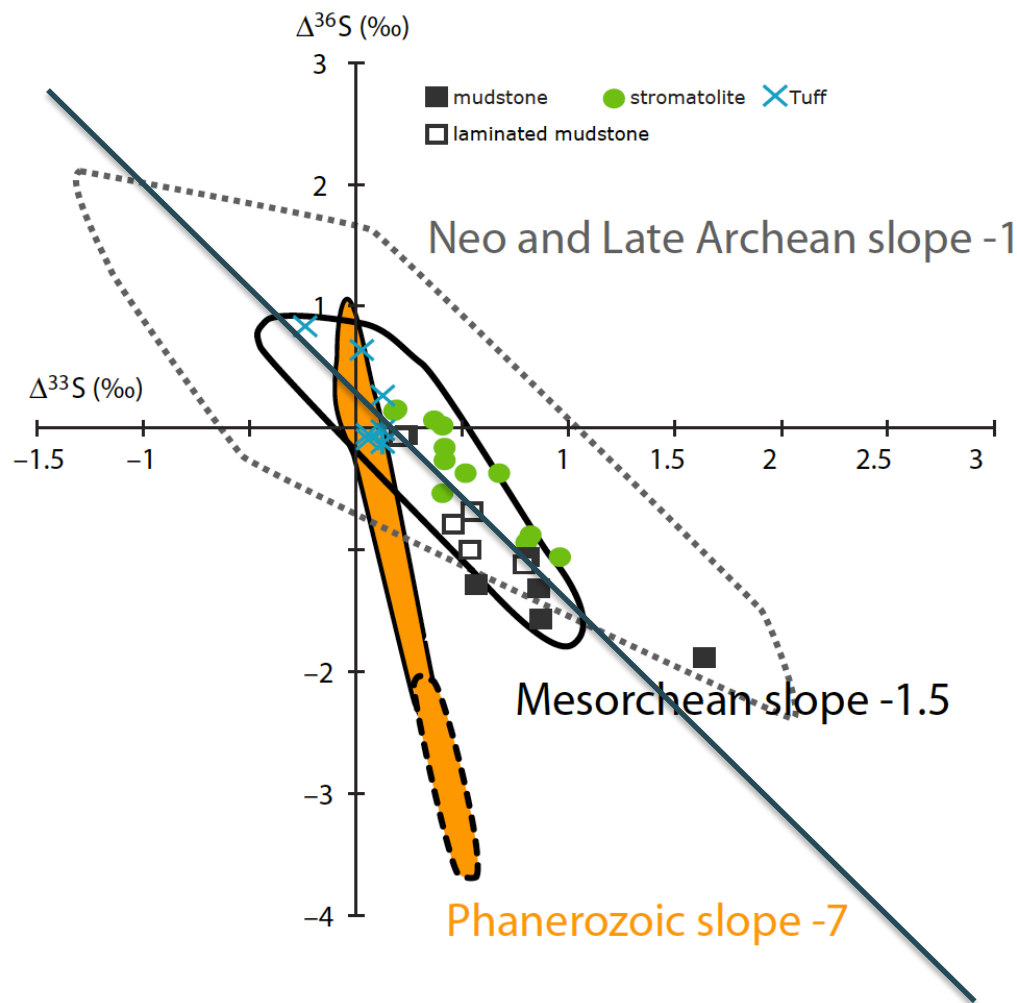
Thomazo et al, 2009



$\Delta^{33}\text{S}_{\text{py}}$ between -0.3 and 1.6 ‰
Sizeable ($> \pm 0.2$ ‰)
BUT LOW MIF-S

1/ 2.73 Ga Tumbiana Formation

Thomazo et al, 2009



$\Delta^{33}\text{S}_{\text{py}}$ between -0.3 and 1.6 ‰

Sizeable ($> \pm 0.2$ ‰)

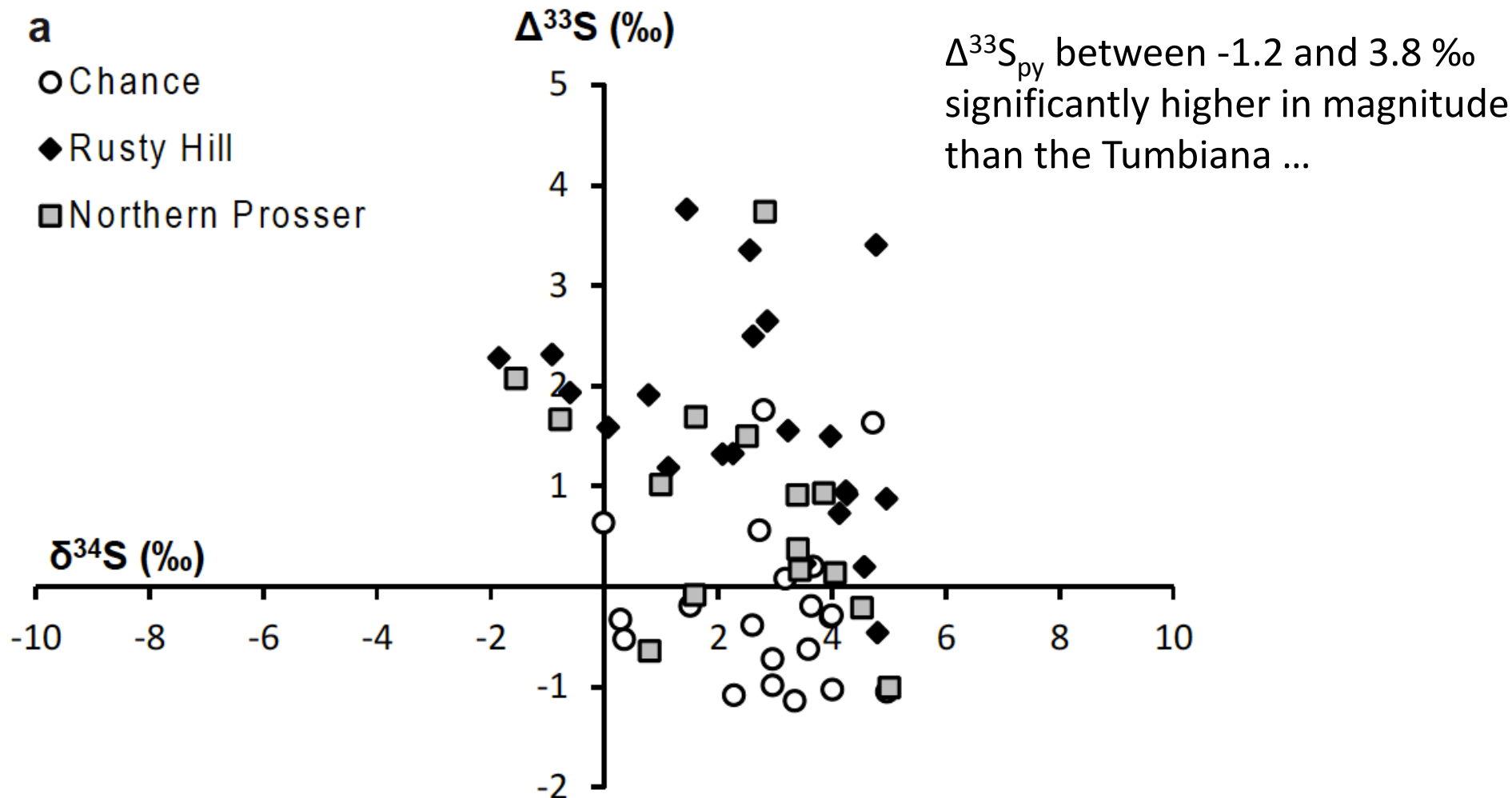
BUT LOW MIF-S

$\Delta^{36}\text{S}/\Delta^{33}\text{S}$ array fall on a typical
Mesoarchean slope of ~ -1.5

➤ Argue for the preservation of
 a Mass Independent signal and
 thus **$\text{O}_2 < 10^{-5}$ PAL**

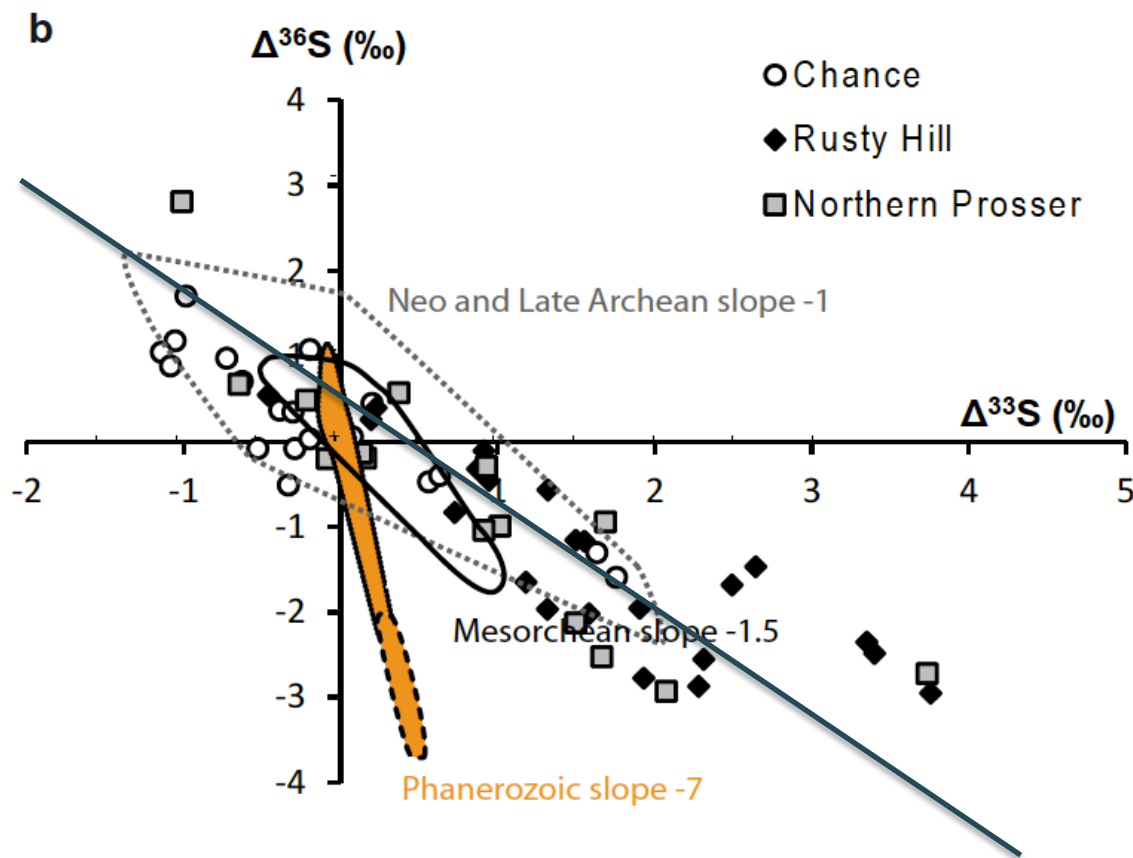
2/ 2.71 Ga Kidd Creek Argillite layers

Kurzweill et al, 2013



2/ 2.71 Ga Kidd Creek Argillite layers

Kurzweill et al, 2013



$\Delta^{33}\text{S}_{\text{py}}$ between -1.2 and 3.8 ‰
significantly higher in magnitude
than the Tumbiana ...

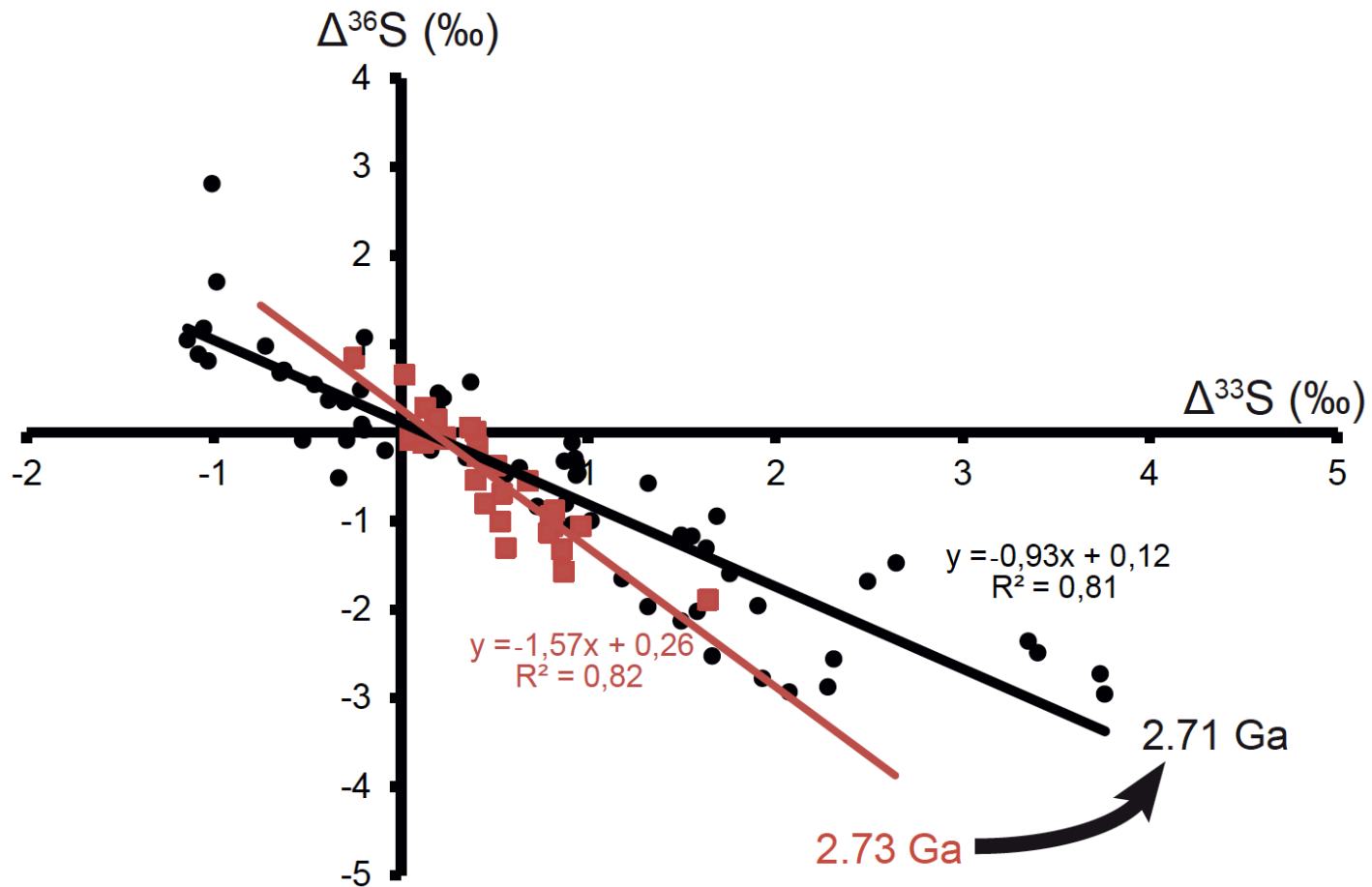
$\Delta^{36}\text{S}/\Delta^{33}\text{S}$ array fall on a typical
Latearchean slope of ~ -1.0

➤ Argue for the preservation of
a Mass Independent signal and
thus **$\text{O}_2 < 10^{-5}$ PAL**

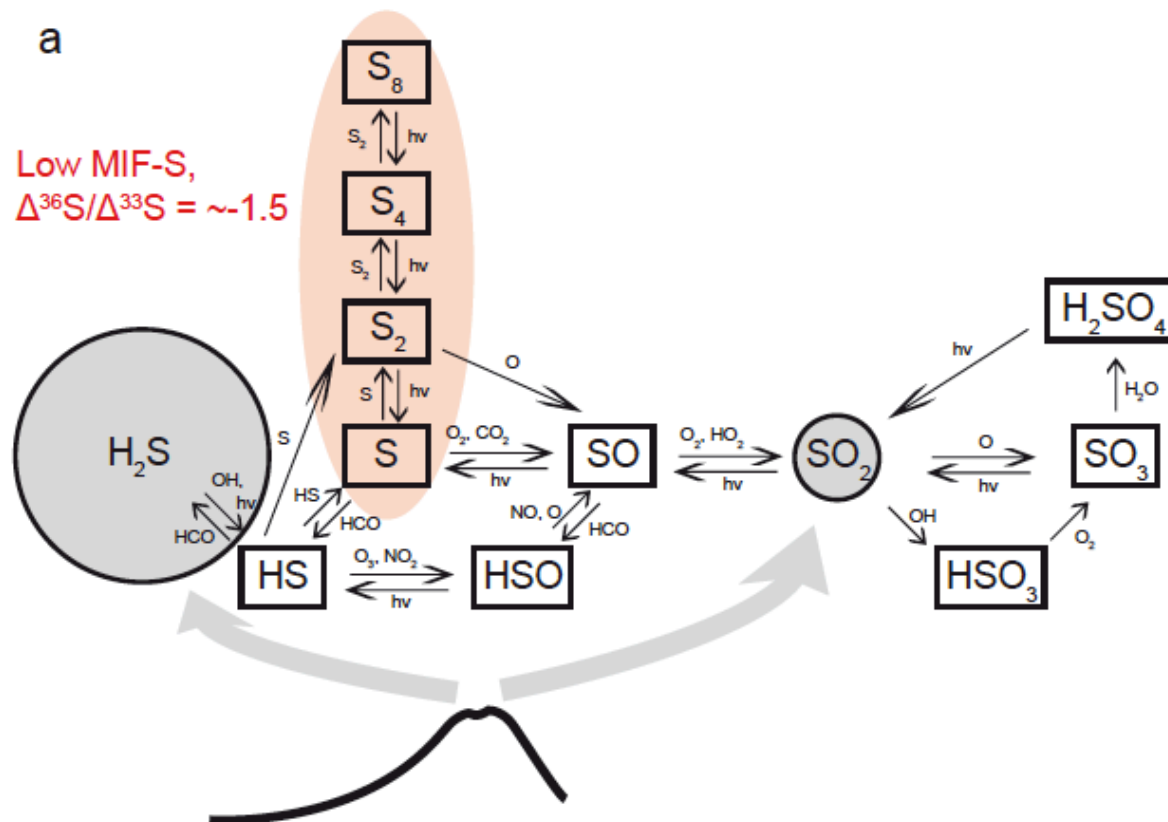
$$Y = -0.93x + 0.12$$

$$R^2 = 0.81$$

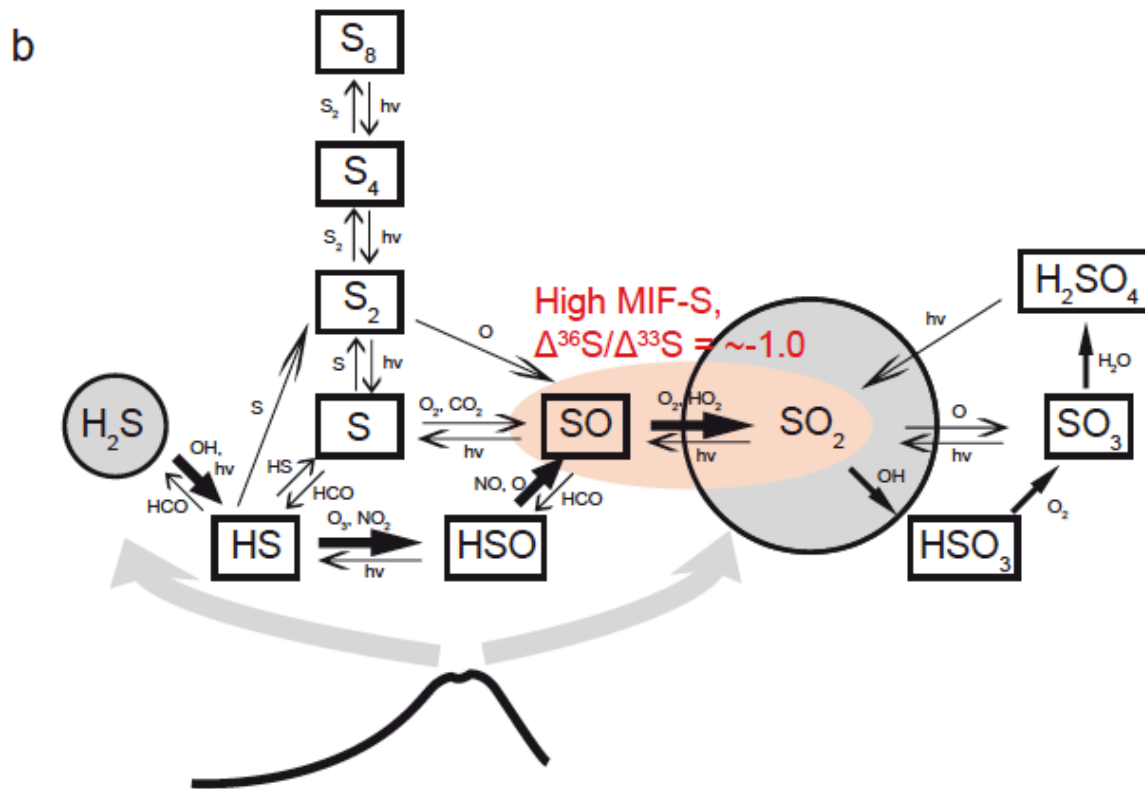
Slope of $\Delta^{36}\text{S}/\Delta^{33}\text{S}$ change between 2.73 and 2.71 Ga !!

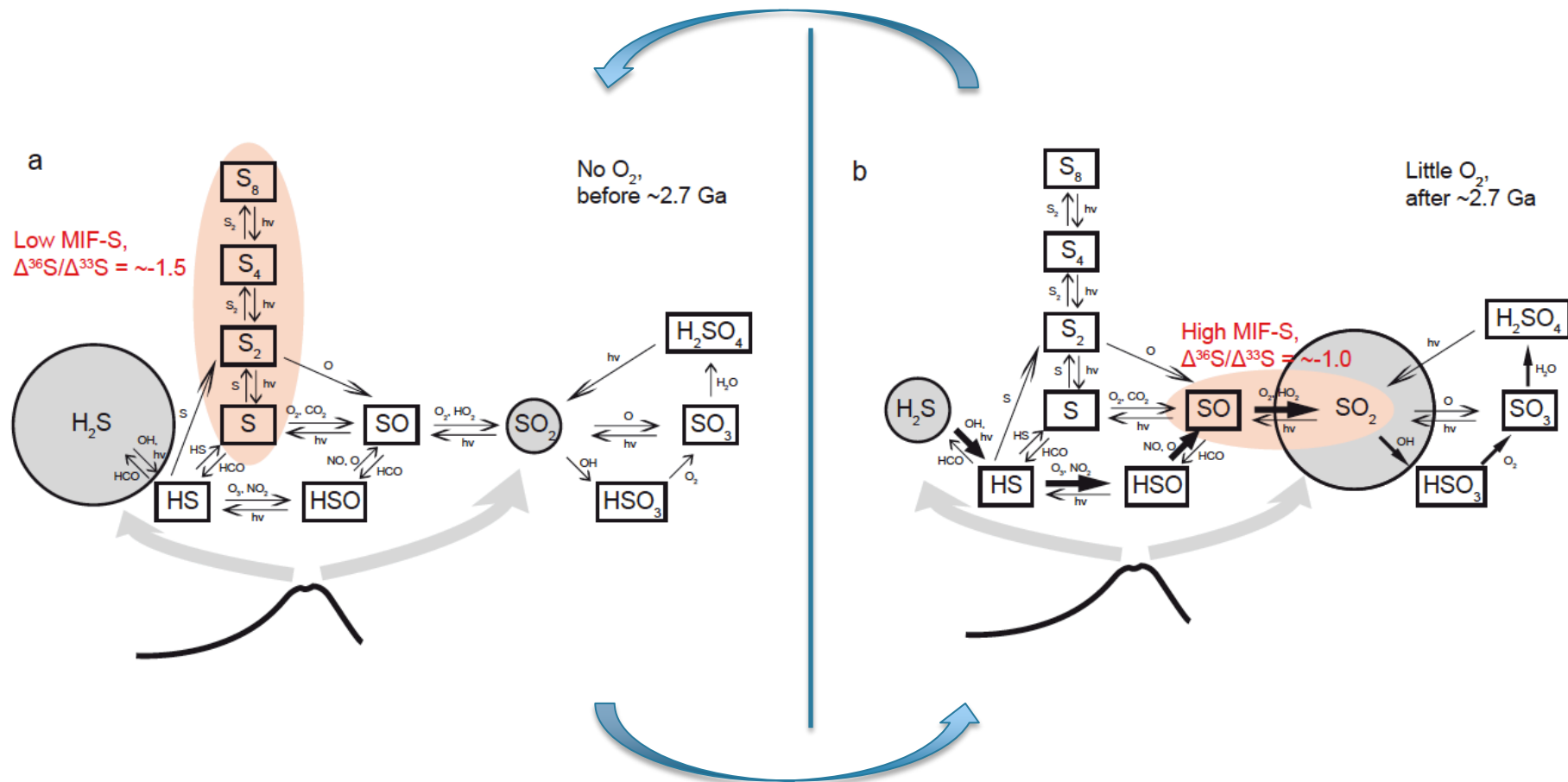


We need some photochemical model ...



We need some photochemical model ...

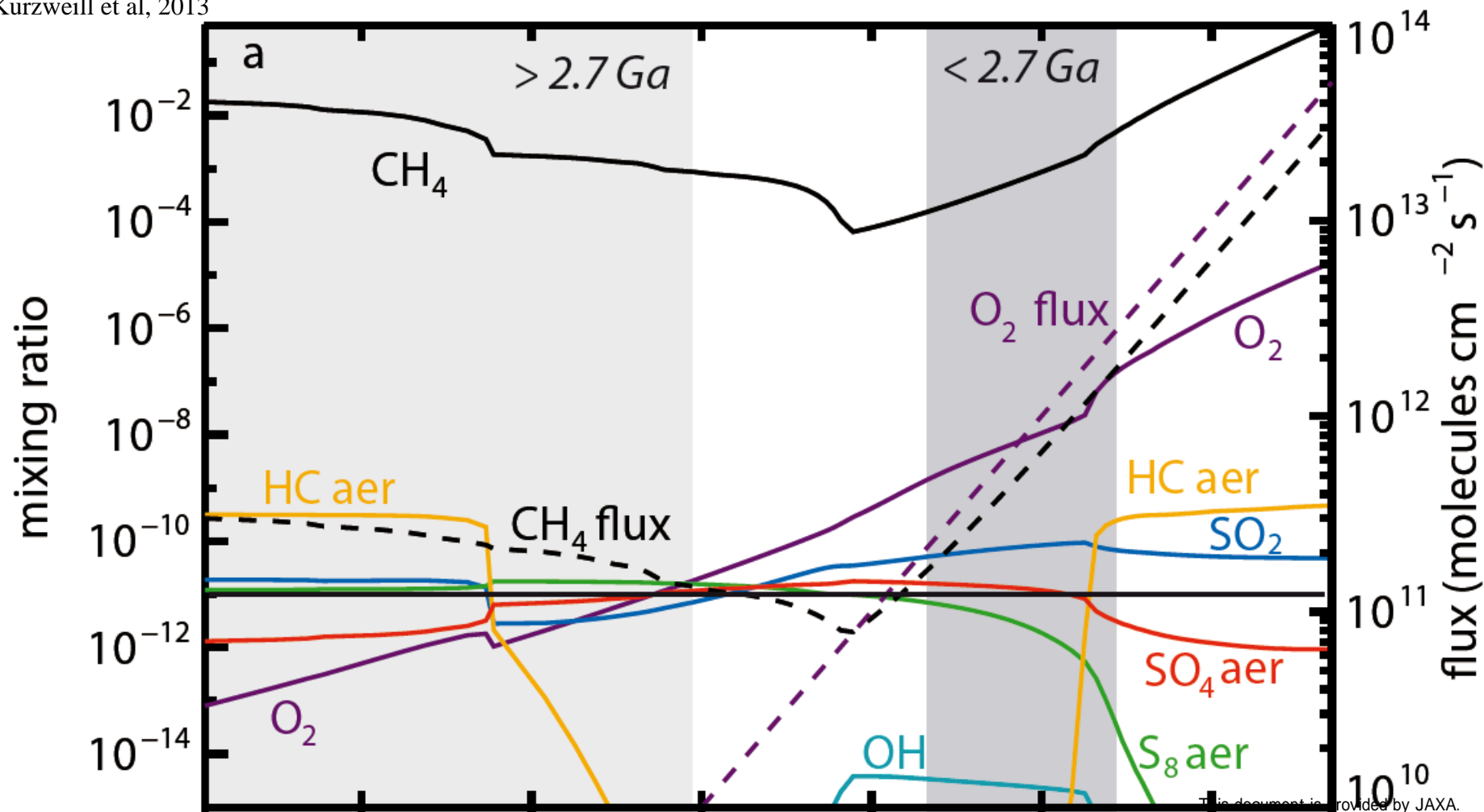


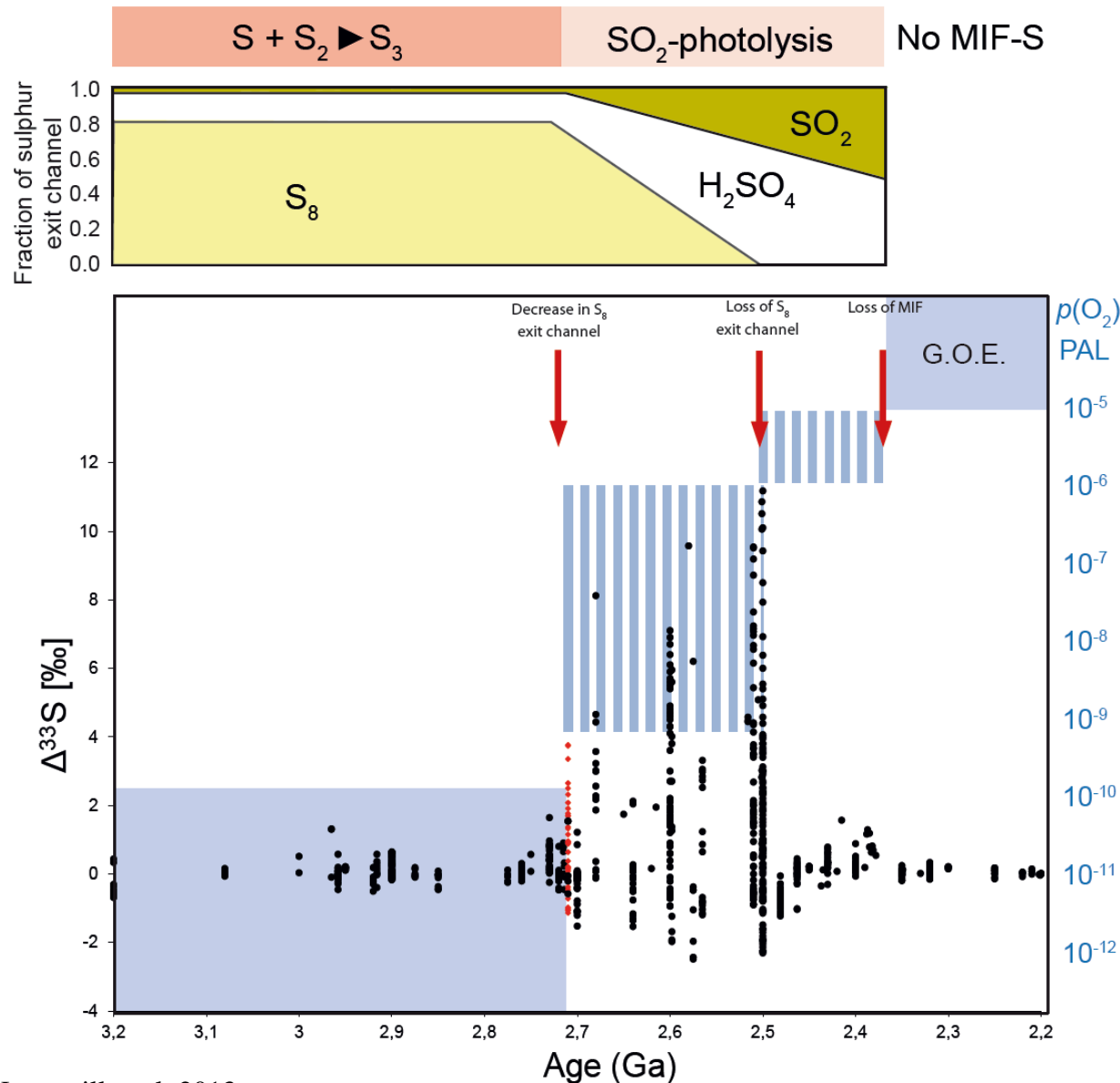


Depends on O₂ mixing ratios

S_8 exit channel starts to decrease at O_2 mixing ratios $>10^{-9}$ consistent with previous estimates of photosynthetic flux into a predominantly reducing atmosphere (Claire et al., 2006)

Kurzweill et al, 2013

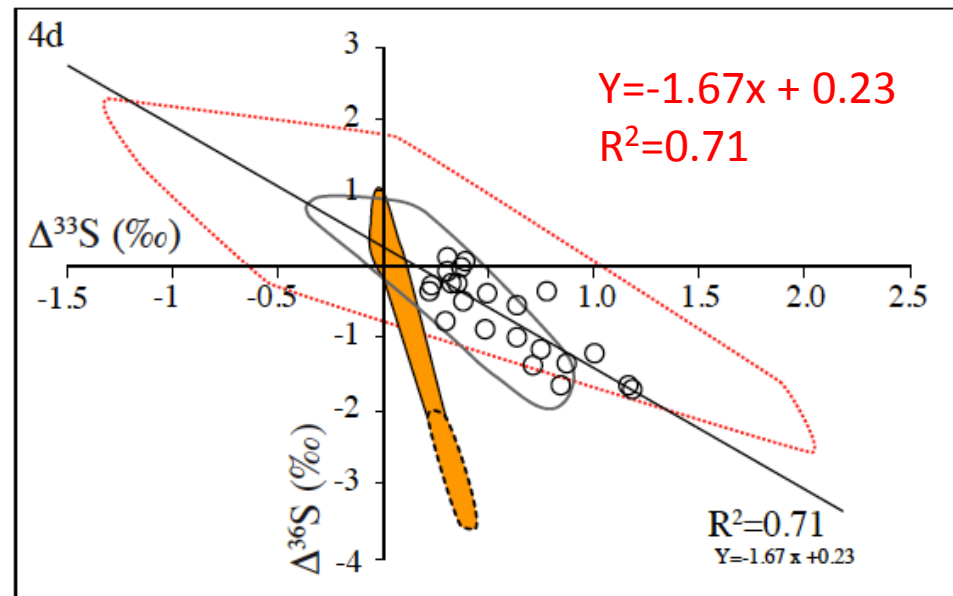
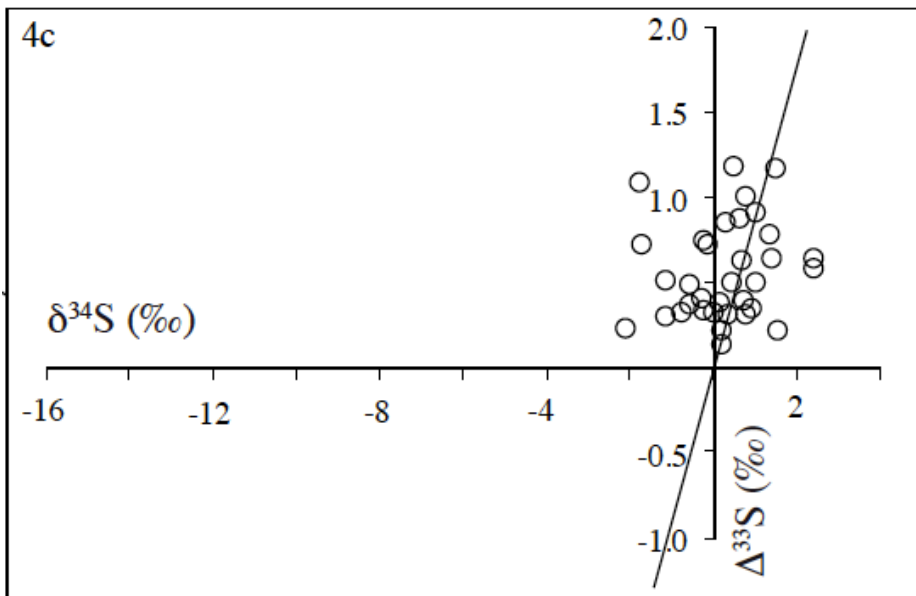




➤ We may have a quantitative model linking MIF-S secular variation and the long term oxygenation of Earth!

However at 2.65 Ga Manjeri and Cheshire formations

Thomazo et al, 2013



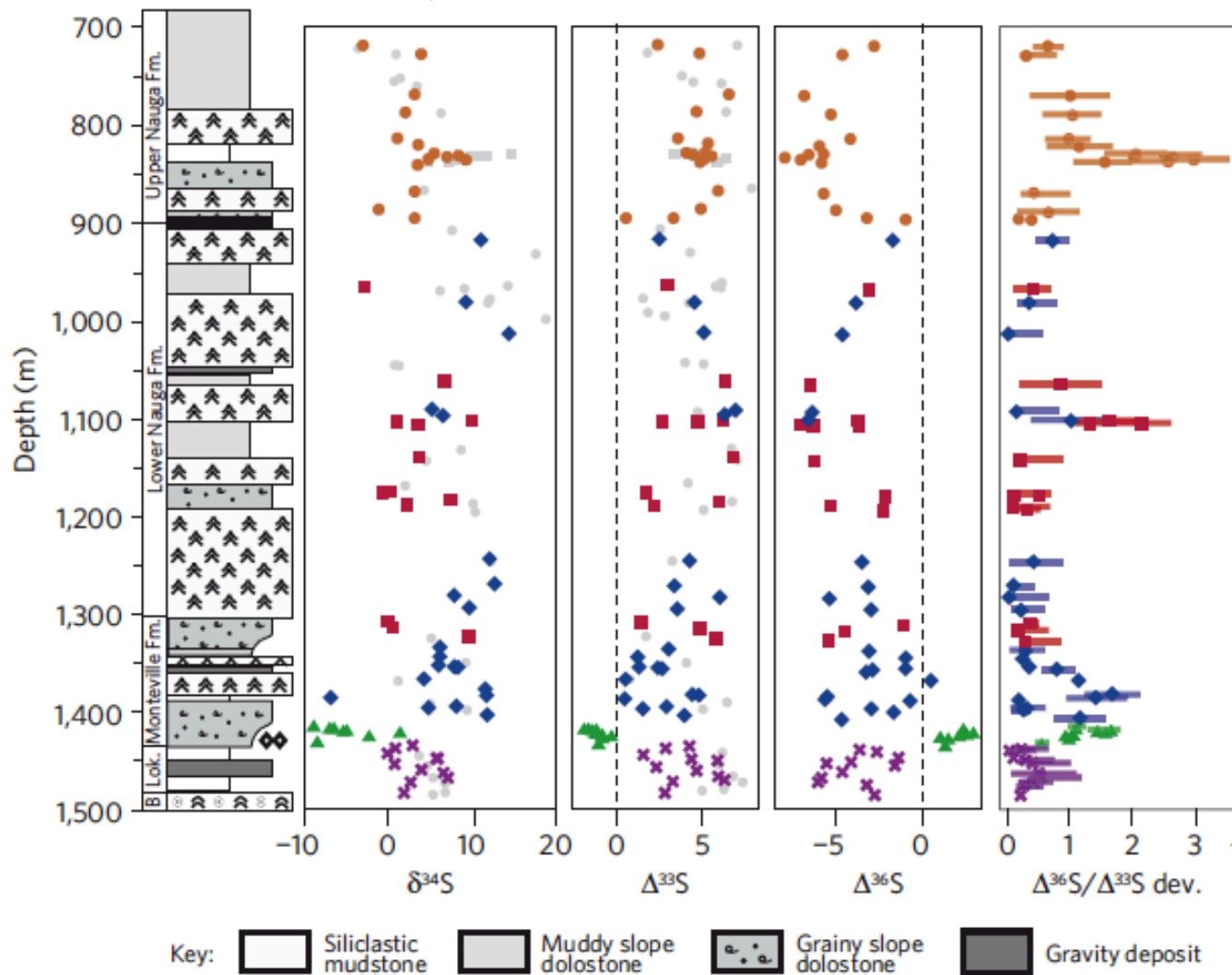
$\Delta^{33}\text{S}_{\text{py}}$ between 0.14 and 1.17 ‰ = Sizeable ($> \pm 0.2$ ‰) BUT LOW MIF-S!!!

$\Delta^{36}\text{S}/\Delta^{33}\text{S}$ array fall on a typical **Mesoarchean** slope of ~ -1.5 !!!

➤ **Our model is wrong / atmosphere is bi-stable / local conditions ?**

GKF01

Zerkle et al, 2012

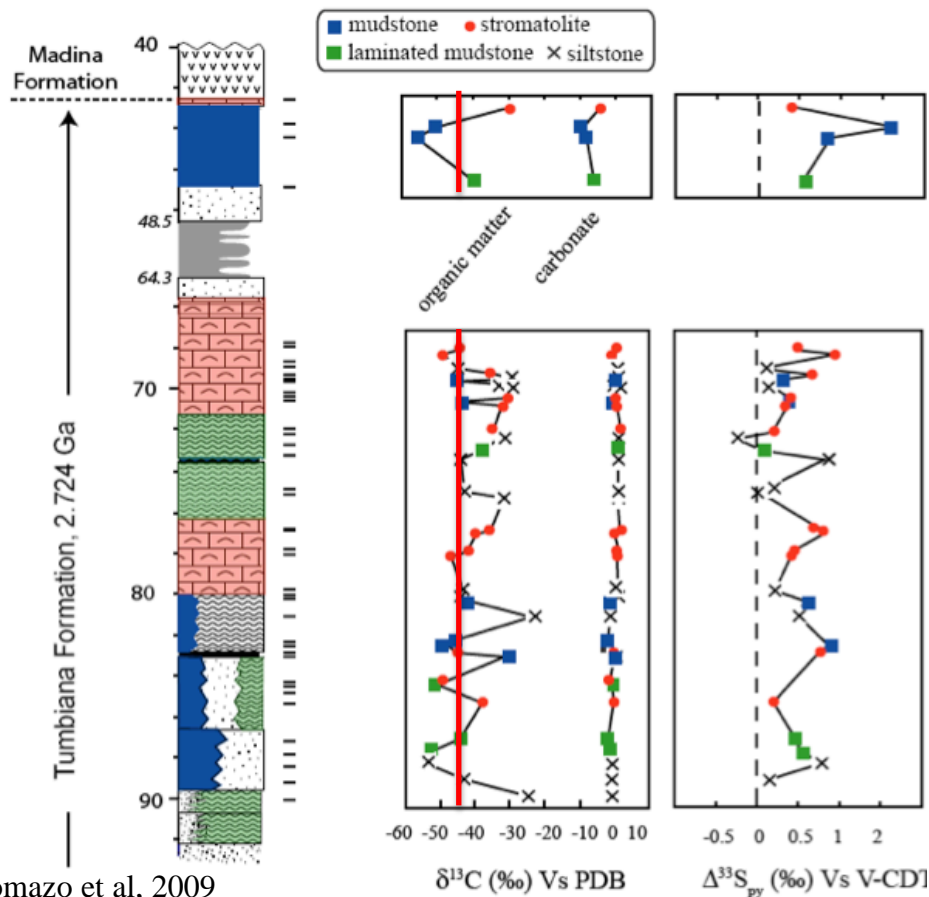


Short term deviation in the $\Delta^{36}\text{S}/\Delta^{33}\text{S}$ are **also observe** in the Campbellrand-Malmani platerofrm deposits (Transvaal Supergroup, South Africa) at around 2.5 Ga.

➤ Our model is wrong / atmosphere is bi-stable / local conditions ?

Neoproterozoic MIF-S oscillations (low MIF-S and $\Delta^{36}\text{S}/\Delta^{33}\text{S}$ array of -1.5) are often linked to $\delta^{13}\text{C}_{\text{org}}$ variations!

In the Tumbiana Formation

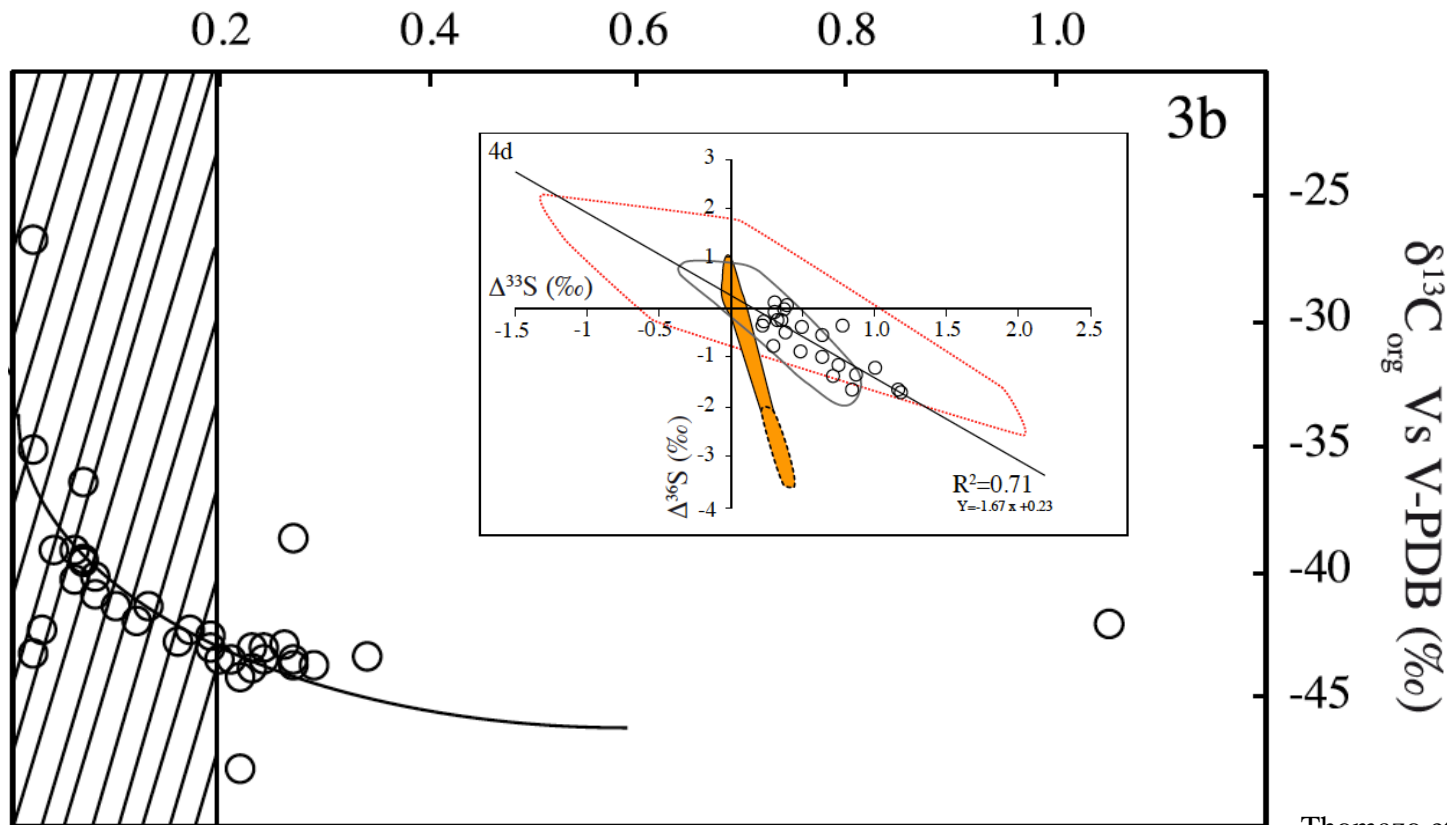


➤ The ~2.72 Tumbiana Fm. record $\delta^{13}\text{C}_{\text{org}}$ range between -55.9 and -23.8 ‰

Neoarchean MIF-S oscillations (low MIF-S and $\Delta^{36}\text{S}/\Delta^{33}\text{S}$ array of -1.5) are often link to $\delta^{13}\text{C}_{\text{org}}$ variations!

In the Tumbiana Formation
In the Cheshire Formation

➤ The ~2.65 Cheshire Fm.
 record $\delta^{13}\text{C}_{\text{org}}$ of $-41.3 \pm 3.5\text{‰}$



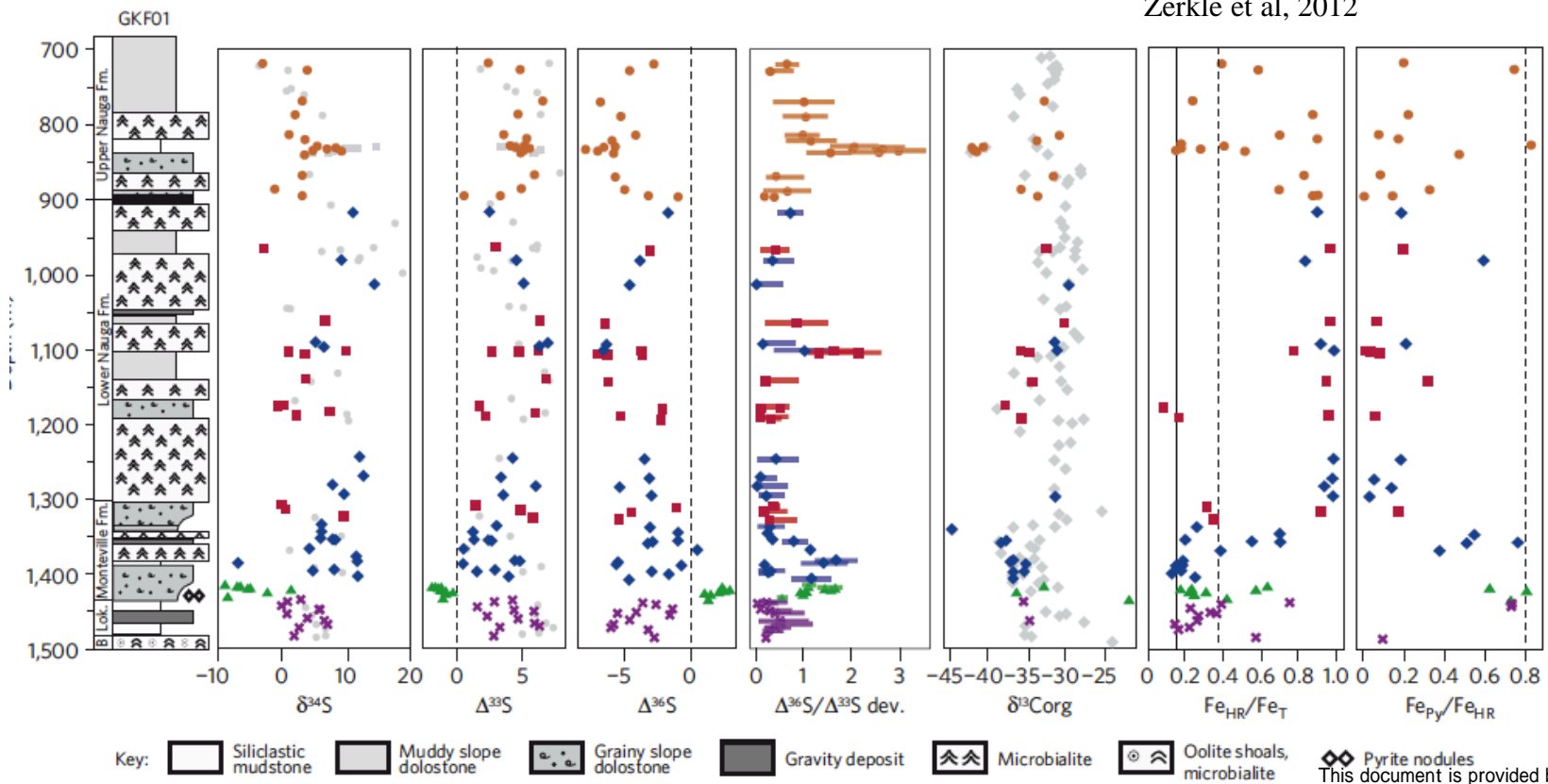
Neoarchean MIF-S oscillations (low MIF-S and $\Delta^{36}\text{S}/\Delta^{33}\text{S}$ array of -1.5) are often link to $\delta^{13}\text{C}_{\text{org}}$ variations!

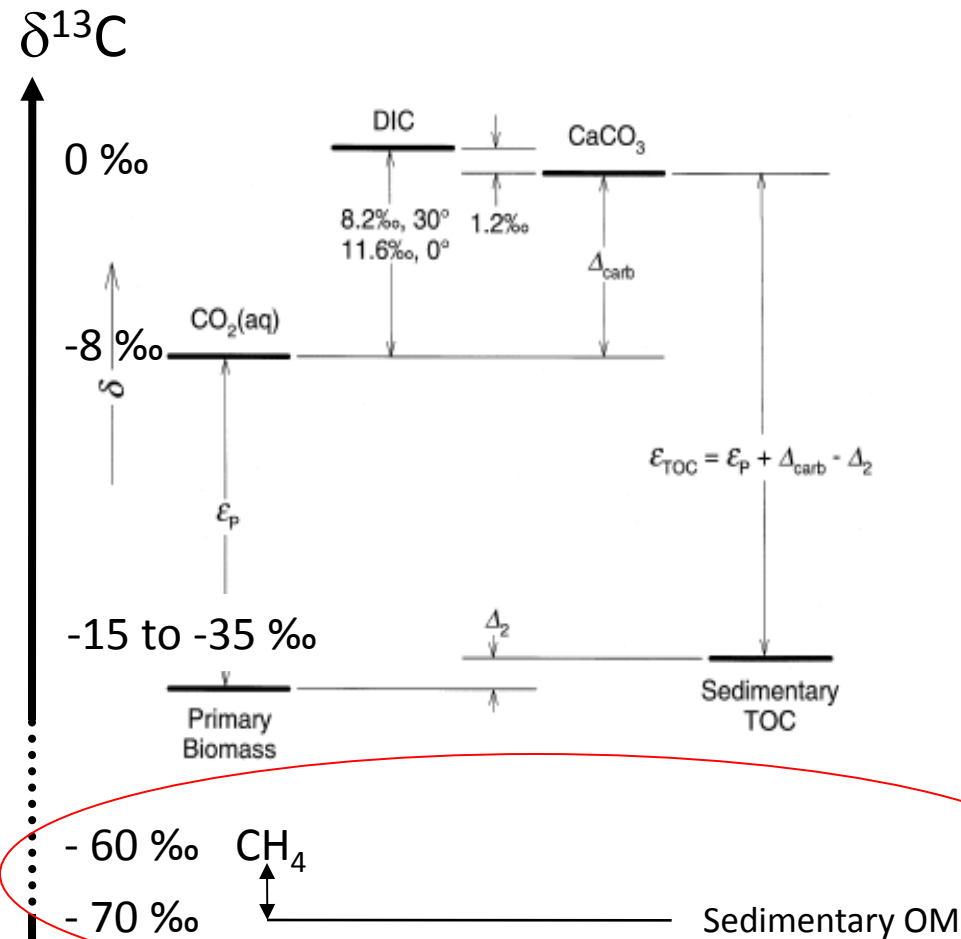
In the Tumbiana Formation

In the Cheshire Formation

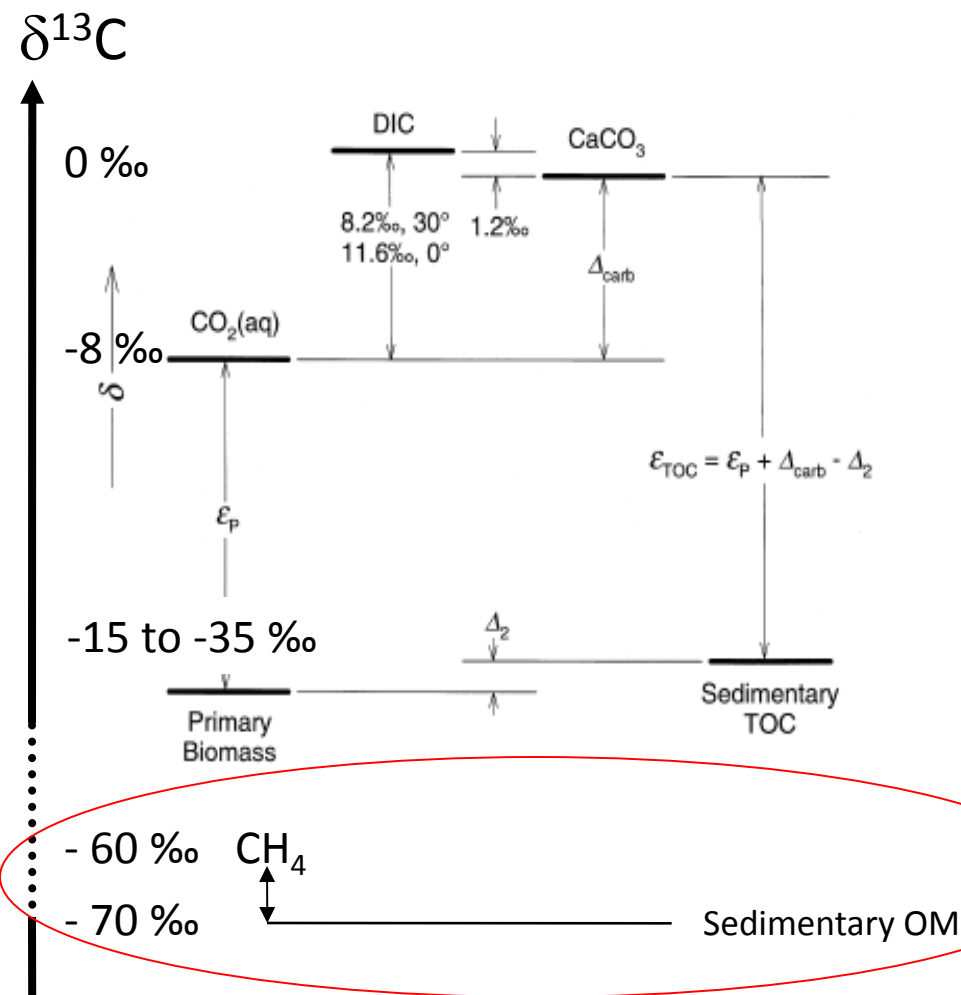
In the Campbellrand-Malmani platerofrm

Zerkle et al, 2012



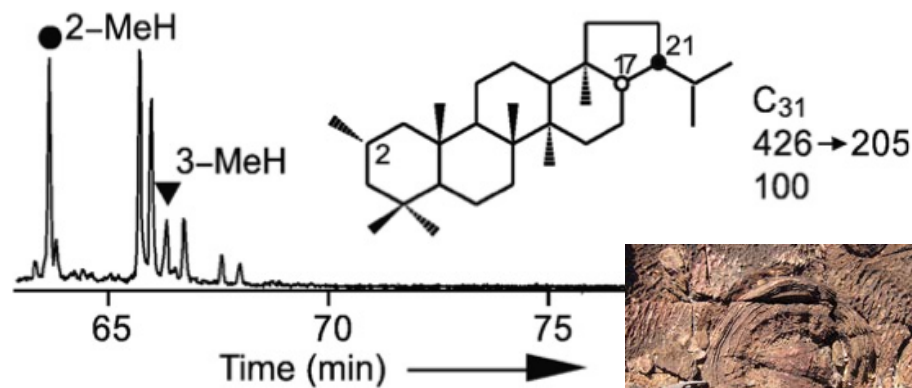


➤ Low $\delta^{13}\text{C}_{\text{org}} < 40\text{‰}$ is interpreted as methane incorporated in organic matter ...



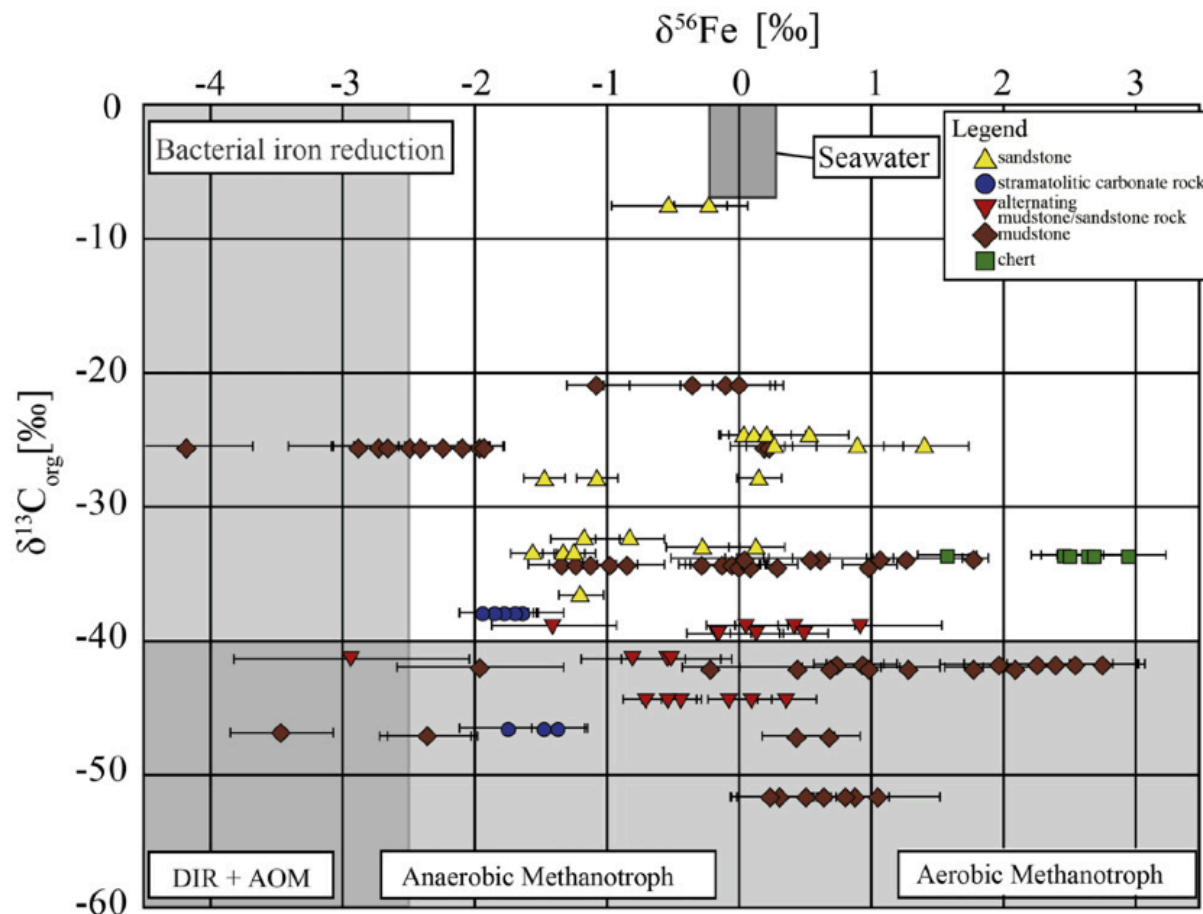
➤ Methane incorporated in organic matter
... by **Microbial oxidation of methane**

2- α methyl hopane

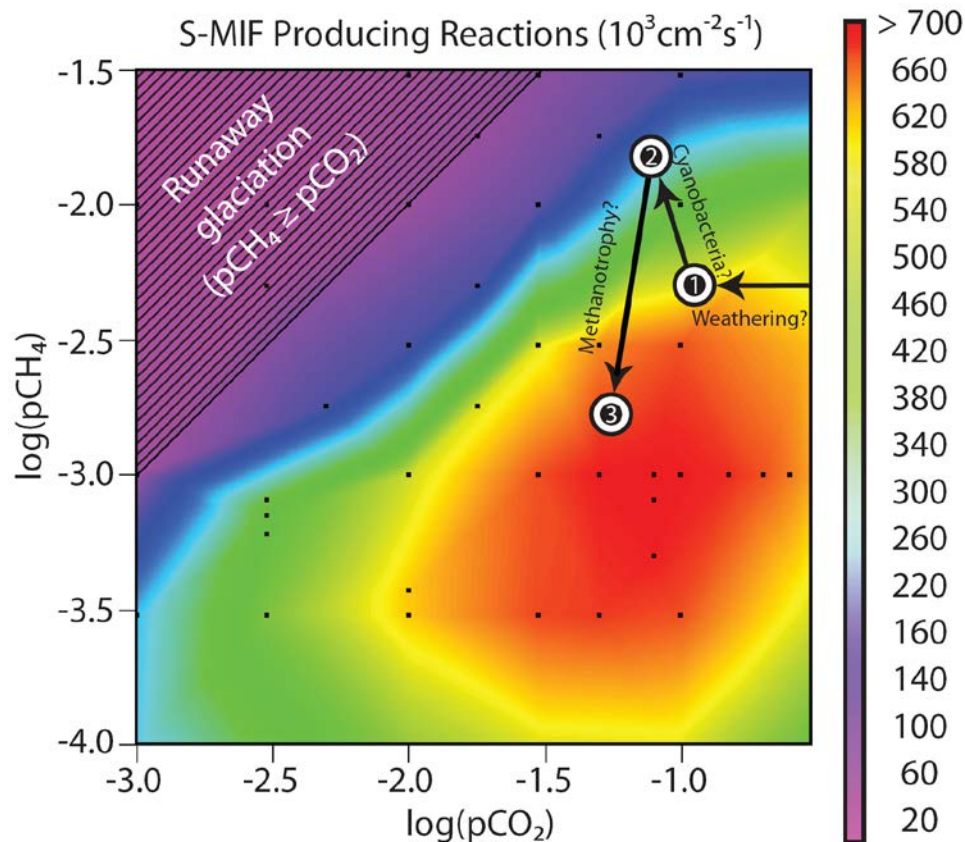


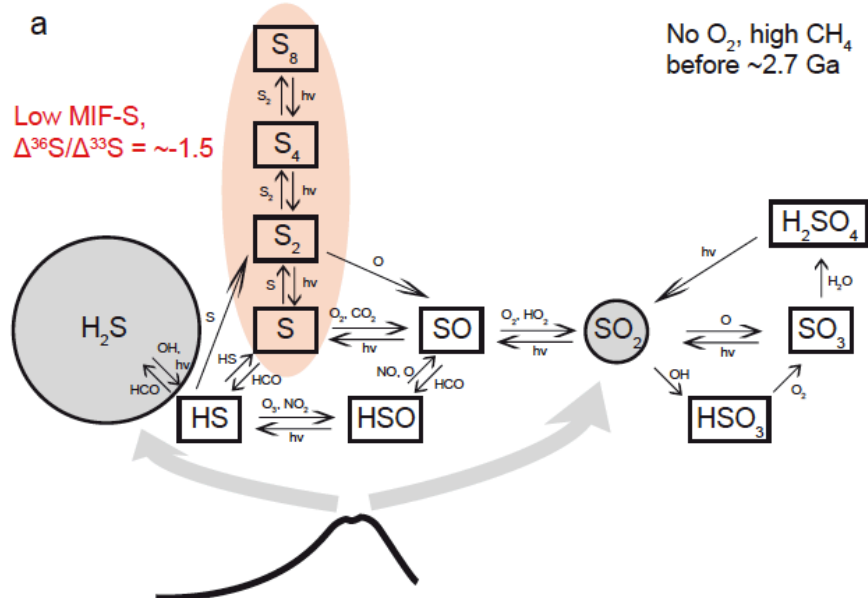
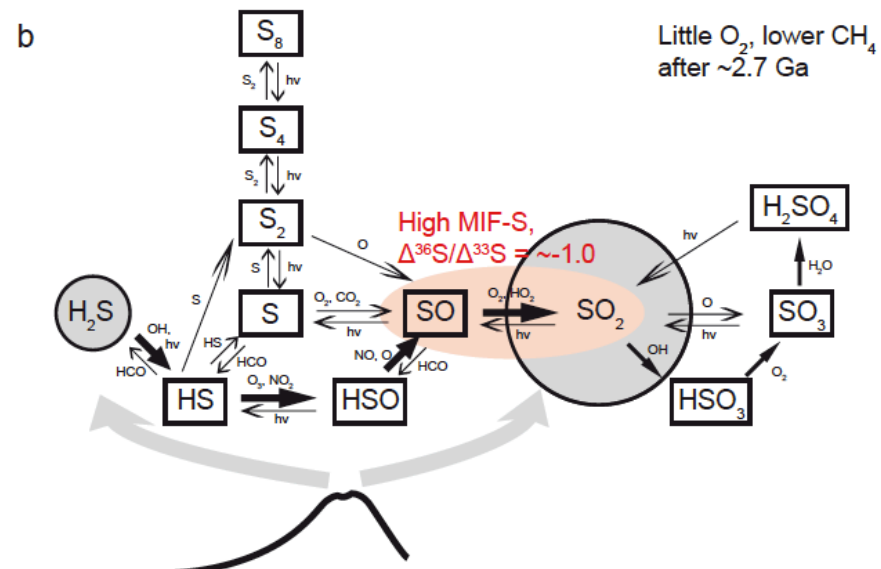
Eigenbrode et al, 2008

- Methanotrophs using **Anaerobic Oxidation of Methane** couple to **iron reduction, sulfate reduction, nitrate reduction ??**



Methanotrophy could have decrease the flux of CH_4 to the atmosphere and modulate the amplitude, sign and reaction rate of sulfur isotopes ...

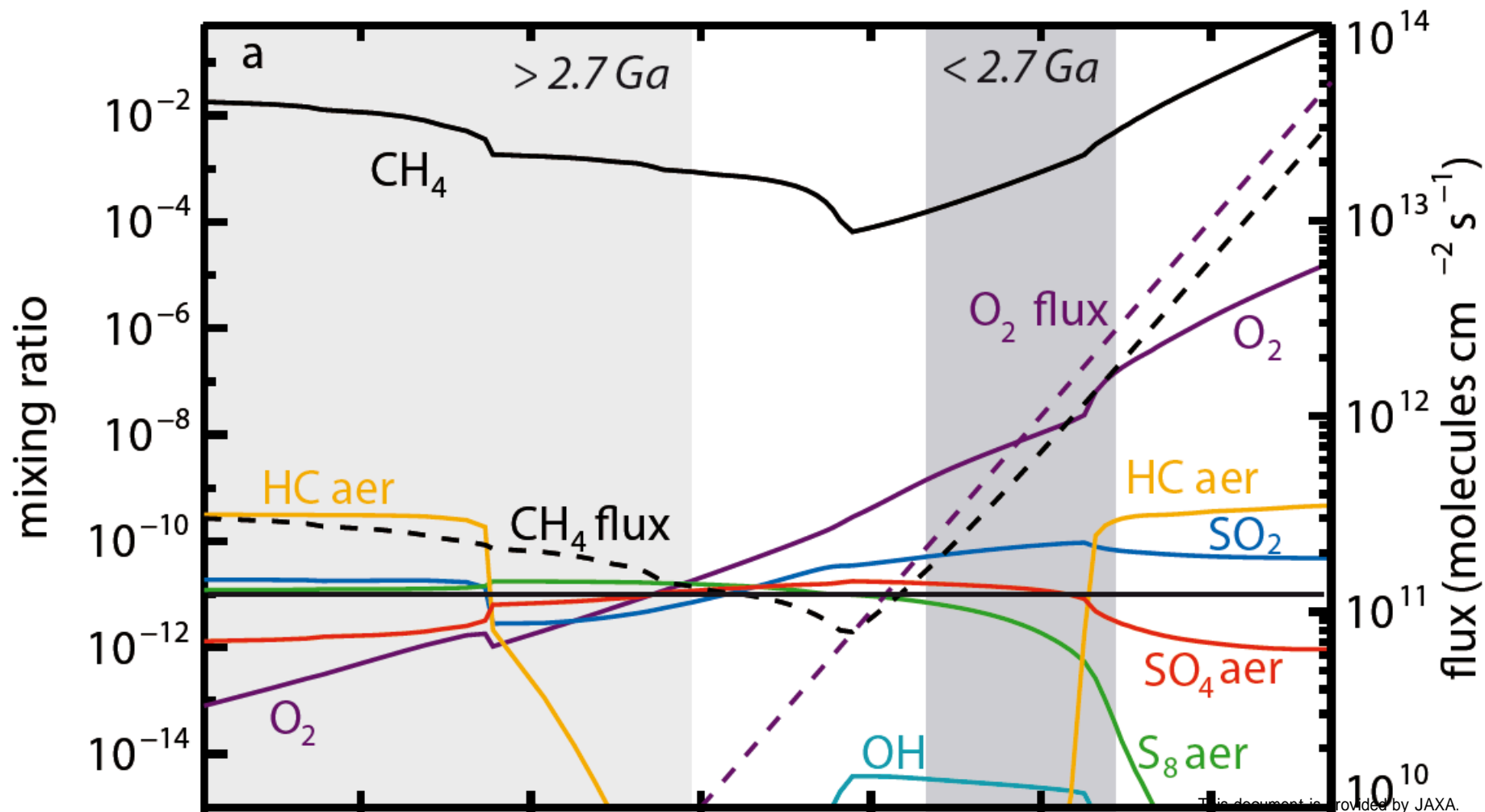


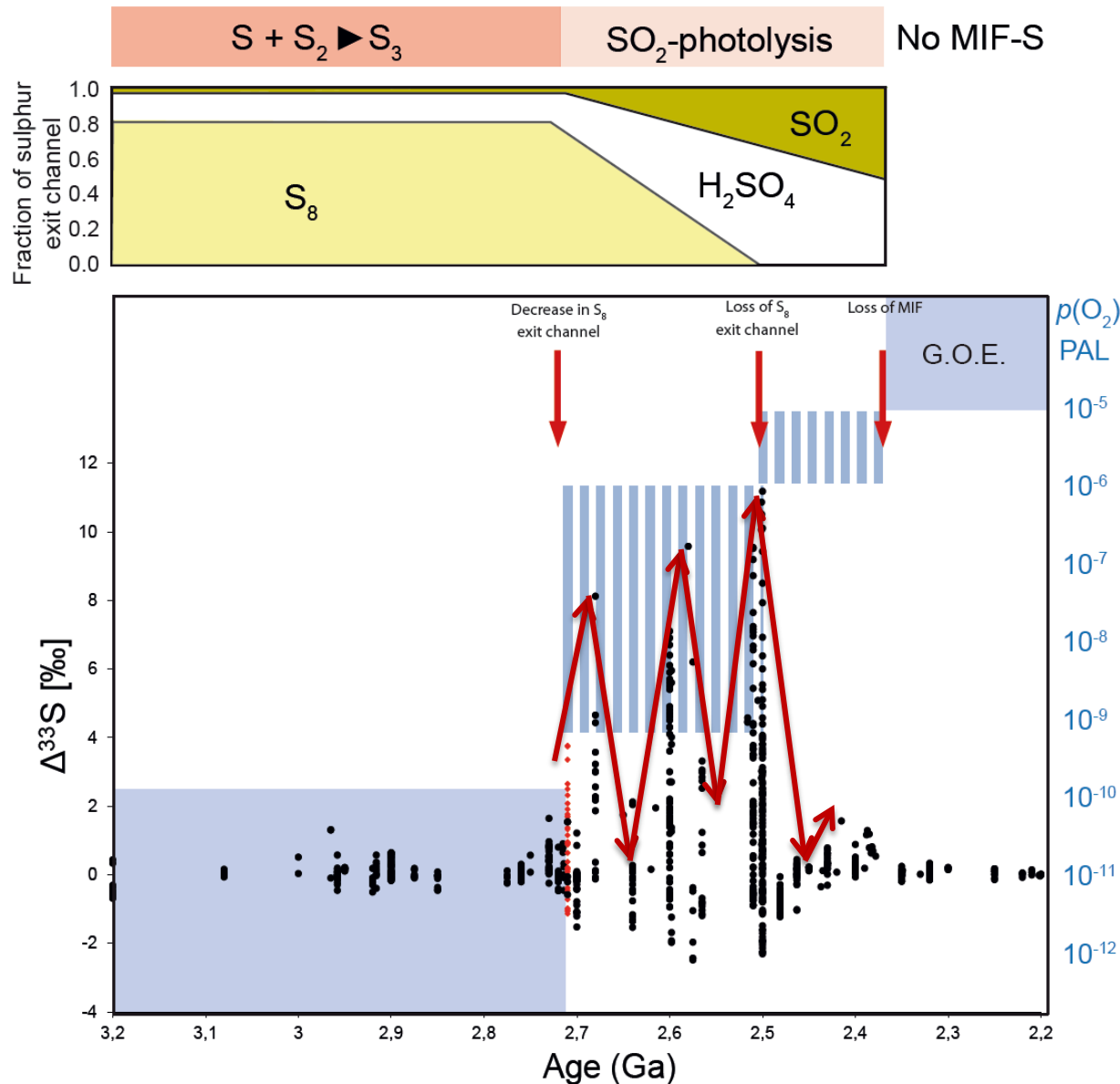
High Atmospheric CH_4 Low Atmospheric CH_4 

Depends on O_2 and $\text{CH}_4:\text{CO}_2$ mixing ratios

S_8 exit channel decrease once the oxygen flux becomes larger than the methane flux ...

Kurzweill et al, 2013





➤ We may have a quantitative model linking MIF-S secular variation and the long term oxygenation of Earth!

➤ Some of the observed variations might be directly link to the oceanic biological methane cycling (in response to a growing oceanic oxidants reservoir : nitrate, sulfate, iron)

Conclusions :

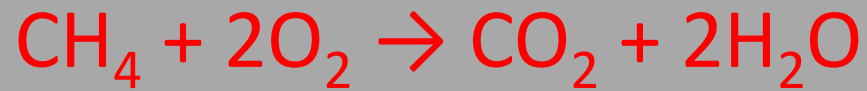
- Model MIF-S secular variations are consistent with a first « whiff of oxygen » at around 2.7 Ga (10^{-11} to 10^{-7} PAL for O_2) in a still largely reducing atmosphere and suggest a strong air–sea gas exchange disequilibrium with respect to O_2
- Some of the observed MIF-S variations might reflect changes in methane flux to the atmosphere regulated by biological process (i.e. the microbial oxidation of methane) ...

BUT our atmospheric model is yet unrealistic / oversimplified, other gases might have been important : H_2S ([Halevy et al, 2010](#)); OCS ([UENO et al. 2009](#)) etc...

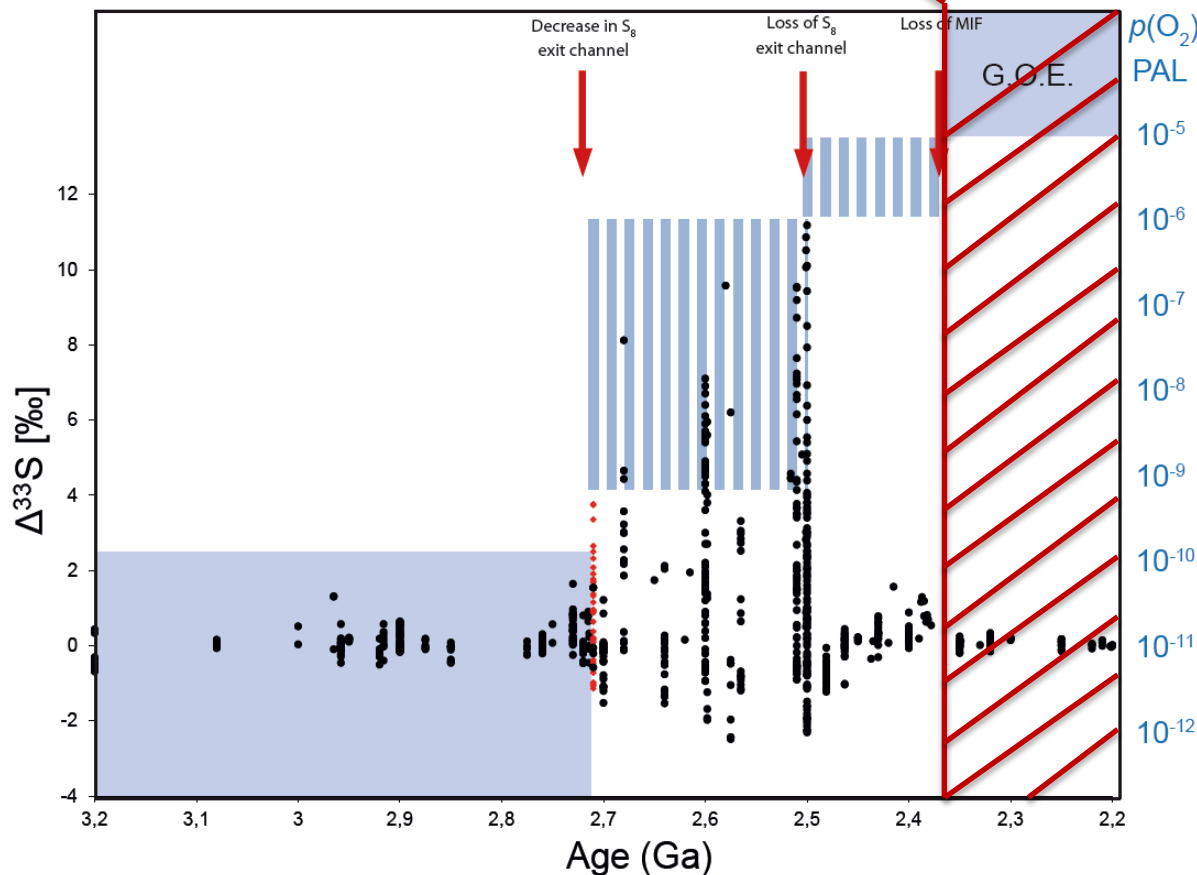
...there is a few 300 Ma of biological and geological feed back loop to decipher ! (sulfur, nitrogen, iron, methane, OM burial, climate) before the GOE!

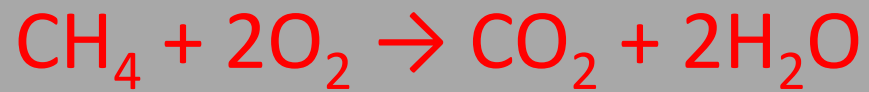
Thanks for listening





➤ After the GOE methane lifetime decrease drastically (atmospheric methane cycle is regulated by abiotic oxidation)





➤ After the GOE methane lifetime decrease drastically (atmospheric methane cycle is regulated by abiotic oxidation)

